

A satellite image of Earth showing a large, dense, white aerosol plume originating from the Indian subcontinent and extending over the Indian Ocean. The plume is visible as a bright, irregular mass against the darker blue of the ocean and the brownish-green of the land. The text is overlaid on the image in a bold, yellow font.

# **Aerosol, Monsoon Rainfall Variability and Climate Change**

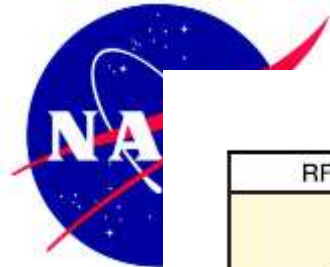
William K. M. Lau  
Laboratory for Atmospheres  
NASA/Goddard Space Flight Center



# Climate Change $\neq$ Global Warming

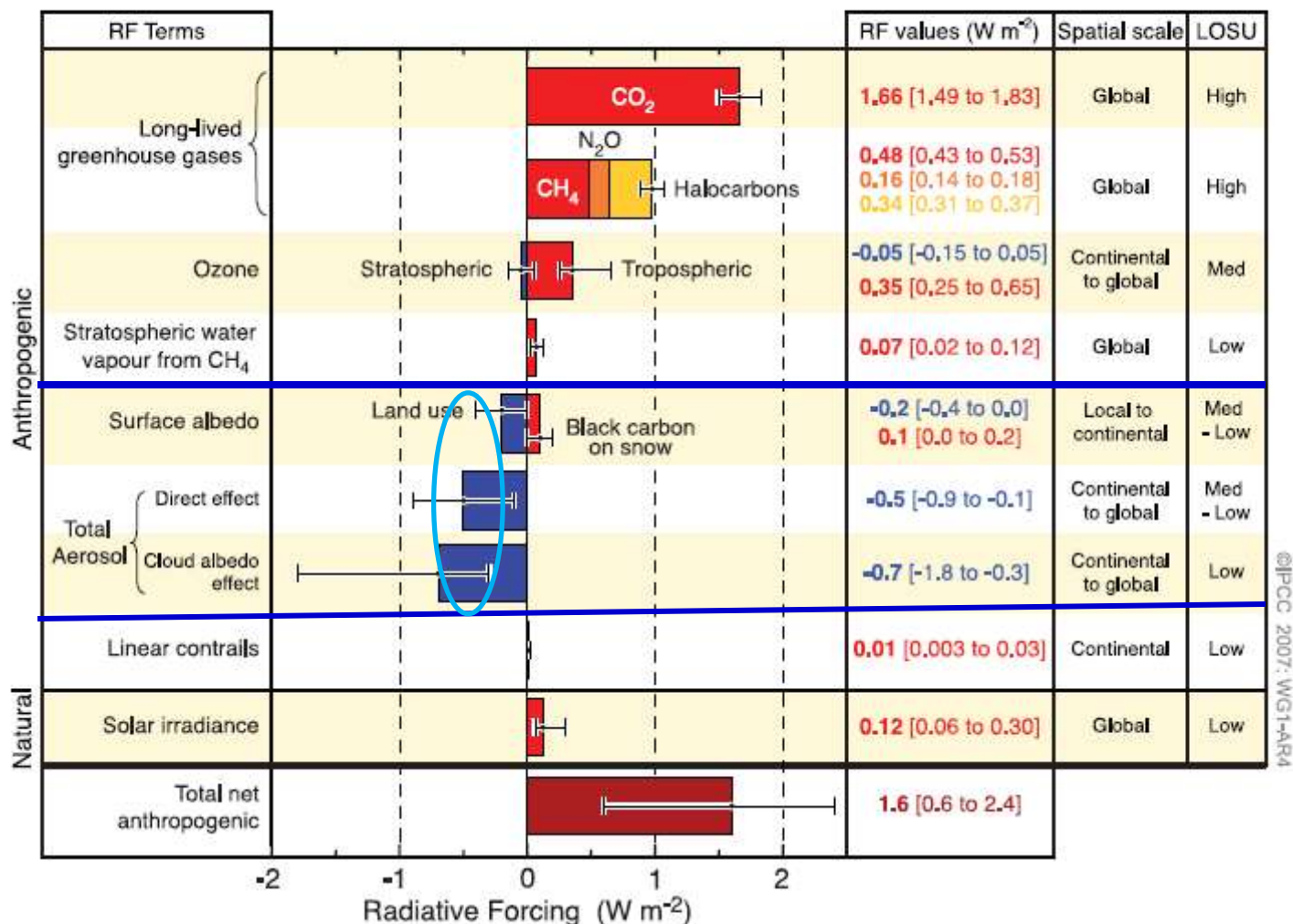
**Climate Change** is a change in the statistical distribution of weather over a period of time that range from season, years, decades to millions of years. It can be a change in the average weather or a change in the distribution of weather events around an average (for example, greater or fewer extreme weather events). Climate change may be limited to a specific region, or an environment, or may occur across the whole Earth System

**Global warming** is a realization of the effects of well-mixed greenhouse gases associated with modern (approx. since the last 1.5 centuries) climate change



## RADIATIVE FORCING COMPONENTS

## Global and TOA



**Figure SPM.2.** Global average radiative forcing (RF) estimates and ranges in 2005 for anthropogenic carbon dioxide ( $\text{CO}_2$ ), methane ( $\text{CH}_4$ ), nitrous oxide ( $\text{N}_2\text{O}$ ) and other important agents and mechanisms, together with the typical geographical extent (spatial scale) of the forcing and the assessed level of scientific understanding (LOSU). The net anthropogenic radiative forcing and its range are also shown. These require summing asymmetric uncertainty estimates from the component terms, and cannot be obtained by simple addition. Additional forcing factors not included here are considered to have a very low LOSU. Volcanic aerosols contribute an additional natural forcing but are not included in this figure due to their episodic nature. The range for linear contrails does not include other possible effects of aviation on cloudiness. {2.9, Figure 2.20}



*".....human influences are significant and involve a diverse range of first-order climate forcings, including , but not limited to the human input of well mixed greenhouse gases (CO<sub>2</sub>, CH<sub>4</sub>, N<sub>2</sub>O etc.) The impacts of these other forcing (aerosols, LULC, urban heat island.....) are critically important in exciting climate feedbacks on regional and local scales, and in modulating the greenhouse warming effects."*

***Climate Change: The Need to Consider Human Forcings Other than Greenhouse Gases, EOS, American Geophysical Union, 2009, Nov.***

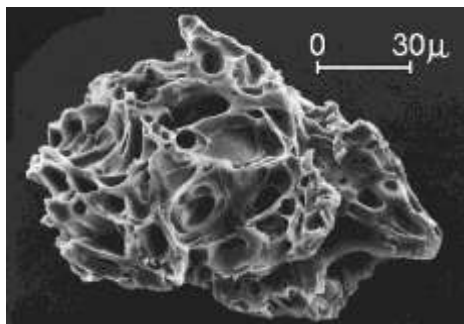
*By Roger Pielke Sr.<sup>1</sup>, Keith Beven<sup>2</sup>, Guy Brasseur<sup>3</sup>, Jack Calvert<sup>3</sup>, Moustafa Chahine<sup>4</sup>, Russ Dickerson<sup>5</sup>, Dara Entekhabi<sup>6</sup>, Efi Foufoula-Georgiou<sup>7</sup>, Hoshin Gupta<sup>8</sup>, Vijay Gupta<sup>1</sup>, Witold Krajewski<sup>9</sup>, E. Philip Krider<sup>8</sup>, **William K. M. Lau<sup>10</sup>**, Jeff McDonnell<sup>11</sup>, William Rossow<sup>12</sup>, John Schaake<sup>13</sup>, James Smith<sup>14</sup>, Soroosh Sorooshian<sup>15</sup>, Eric Wood<sup>14</sup>*



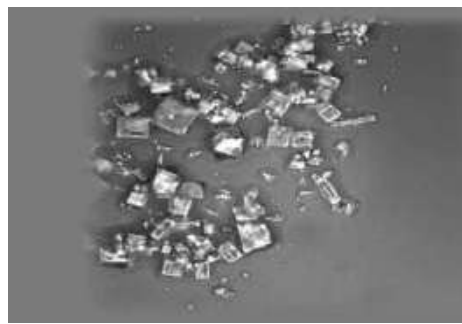
# Definition of aerosol

An **aerosol** is a suspension of fine solid particles or liquid droplets in a gas. Aerosols come from both man-made (industry, motor vehicles), and natural sources (forest fires, oceanic haze, deserts). *Examples are: sulfate, soot (black carbon), organic carbon, dust, sea salt.*

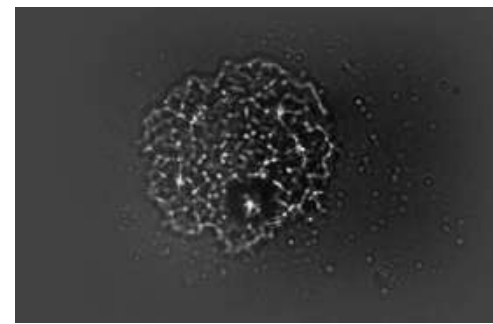
**Soot** is impure carbon particles resulting from incomplete combustion of a hydrocarbon e.g., from internal combustion engines, power plants, heat boilers, waste incinerators, furnaces, fireplaces, slash and burn agriculture, forest fires.



Volcanic ash



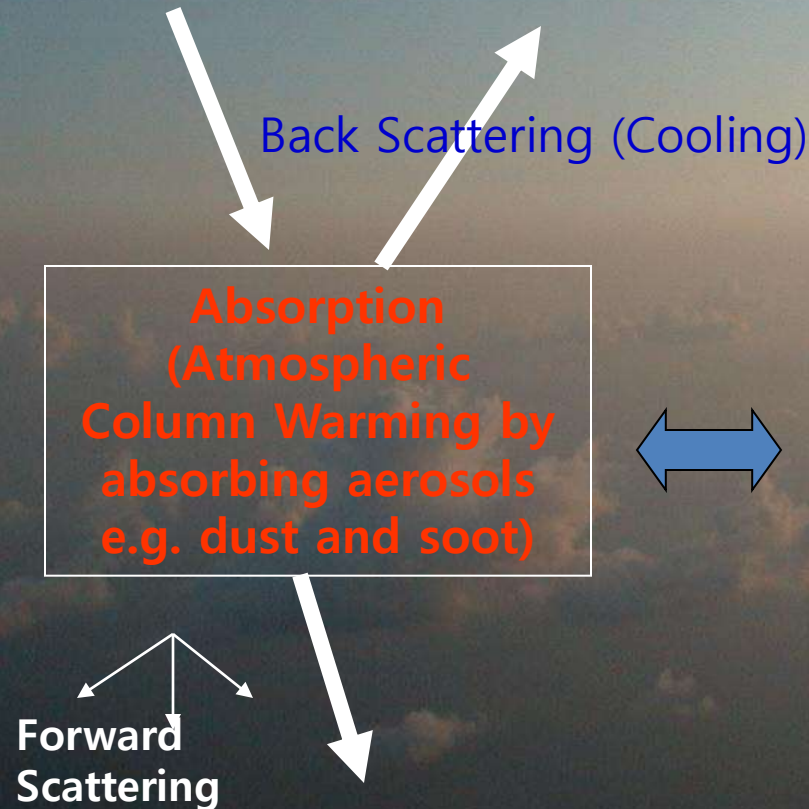
Sea Salt



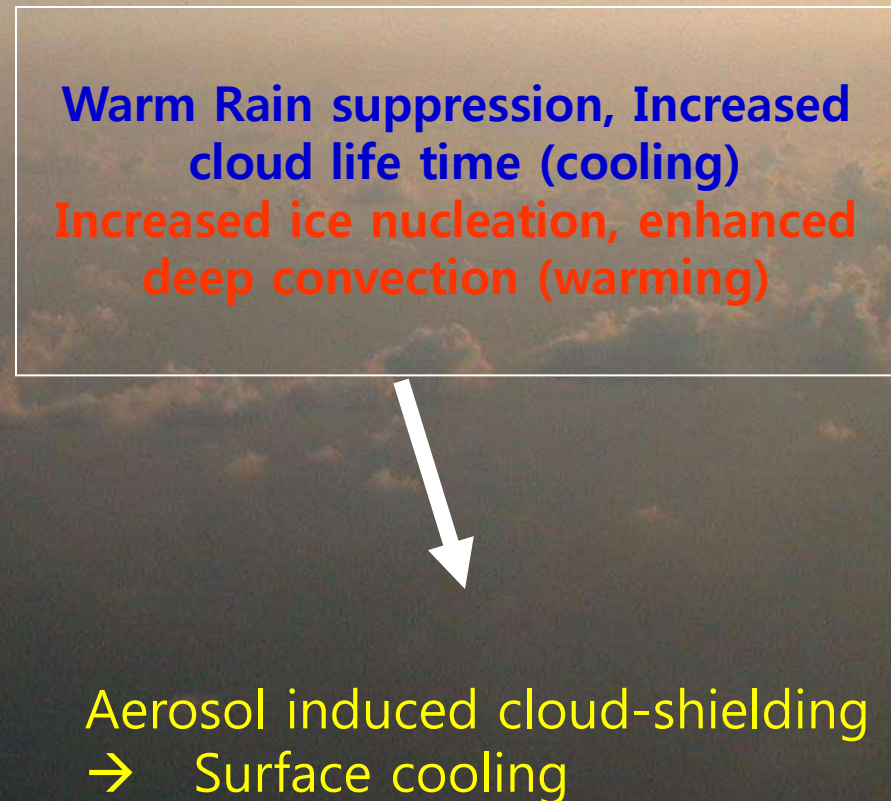
Tractor soot

# Aerosols alter energy balance of the atmosphere-ocean-land system

## Radiation (Direct) Effects

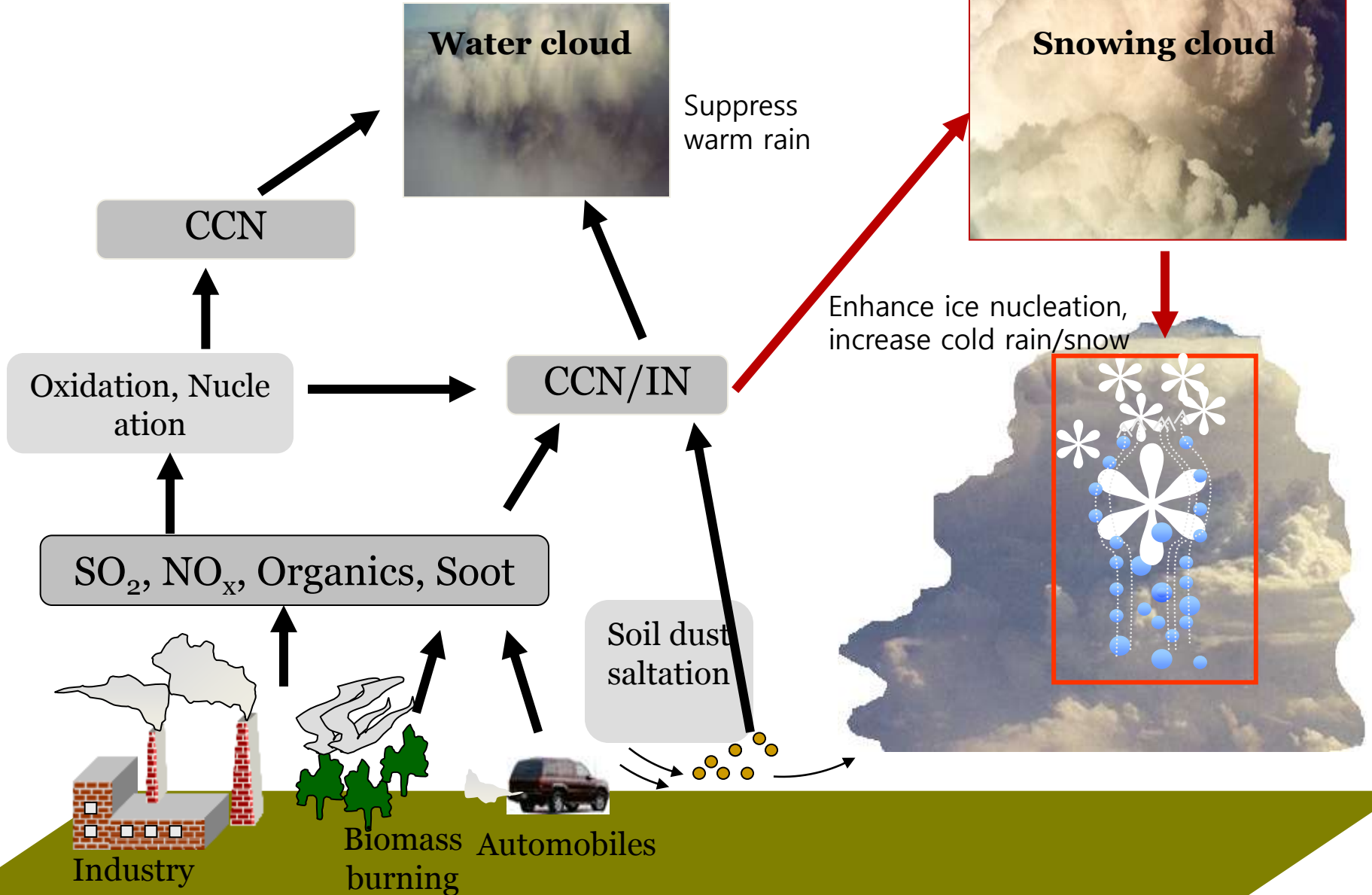


## Microphysics (Indirect) Effects



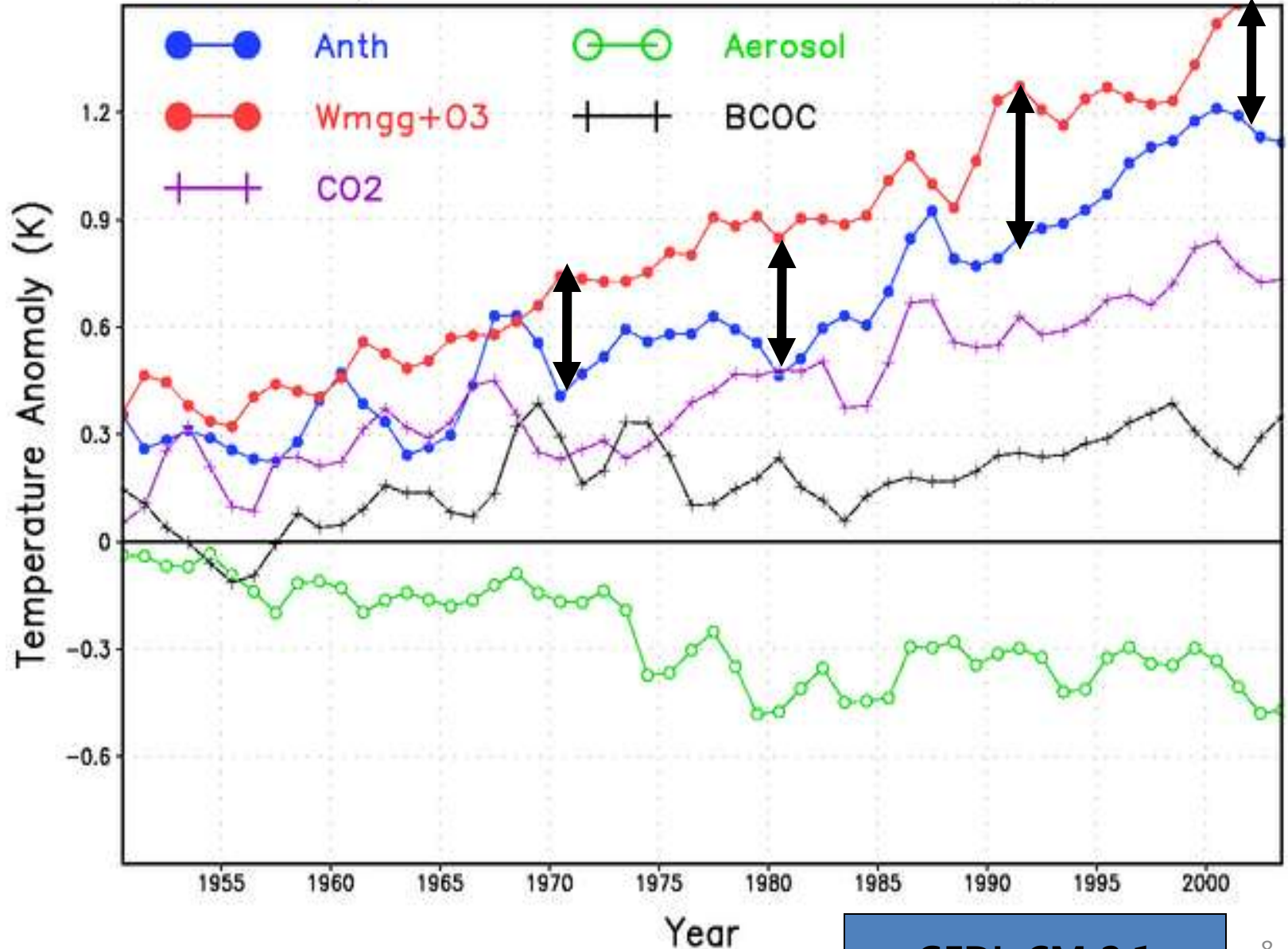


# Microphysics (indirect) effects of aerosols on warm rain and ice-phase rain





# Aerosols mask global warming effects



GFDL CM 2.1



Aerosols have done a lot more than just "offset" the global-mean surface temperature due to **long-Lived Greenhouse Gases**, but also altered regional water cycle and climate in ways that are still not understood, or even known!!

# NASA's Satellites Taking the Pulse of the Earth





# Aerosol-Hotspots

*Aqua Collection-05 Aerosol (Deep Blue & MODIS Composite)*

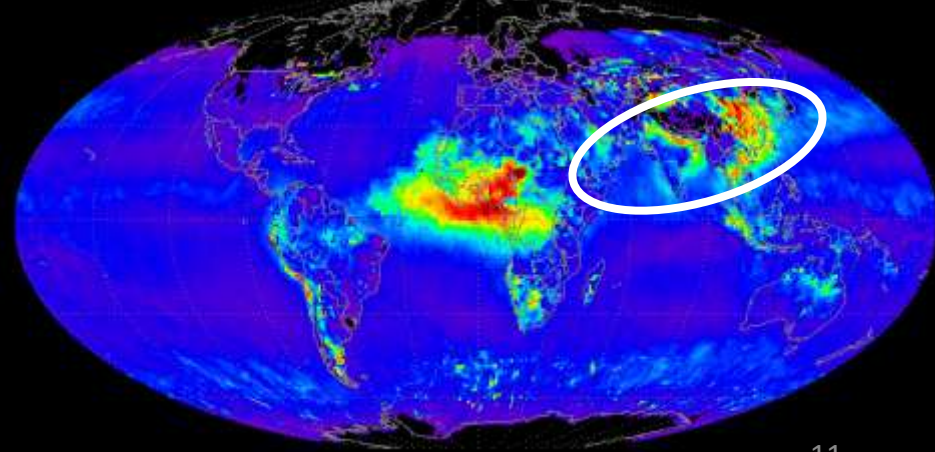
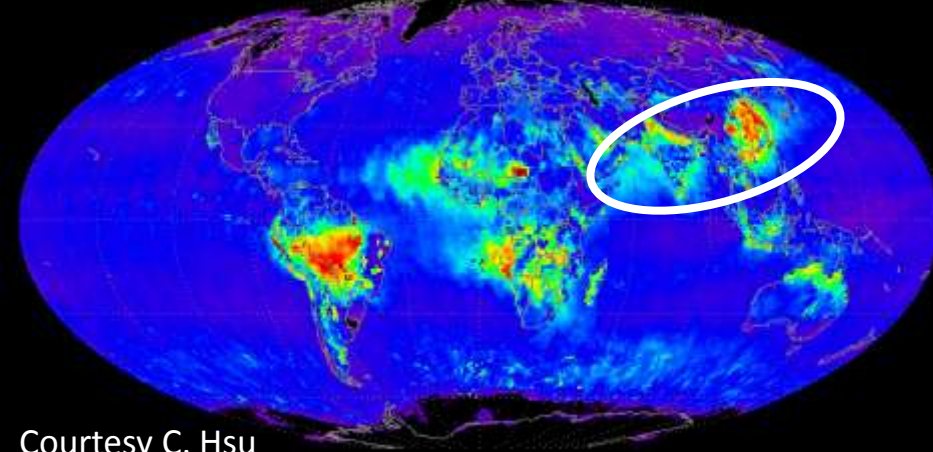
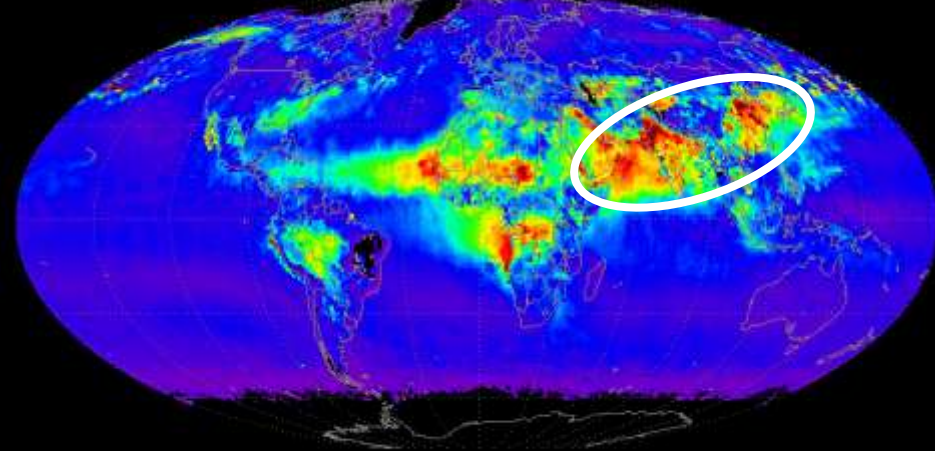
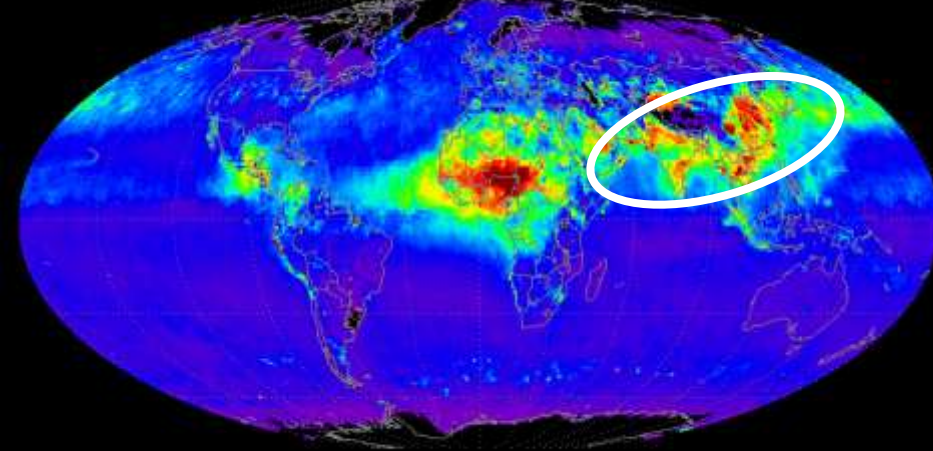
March-April-May

June-July-August

2005

September-October-November

December-January-February



Courtesy C. Hsu



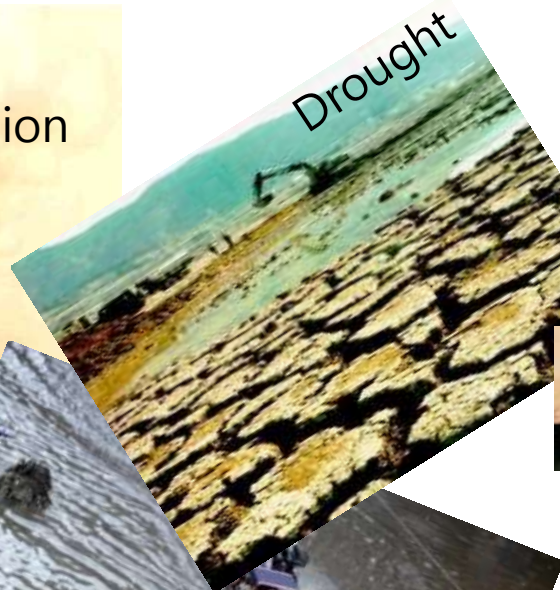
$\tau_{\lambda}(0.55\mu\text{m})$



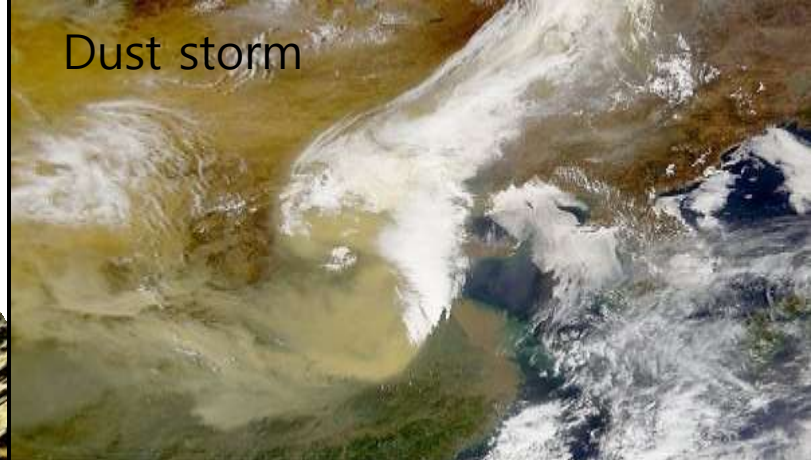
# Climate Change hazards in Asia



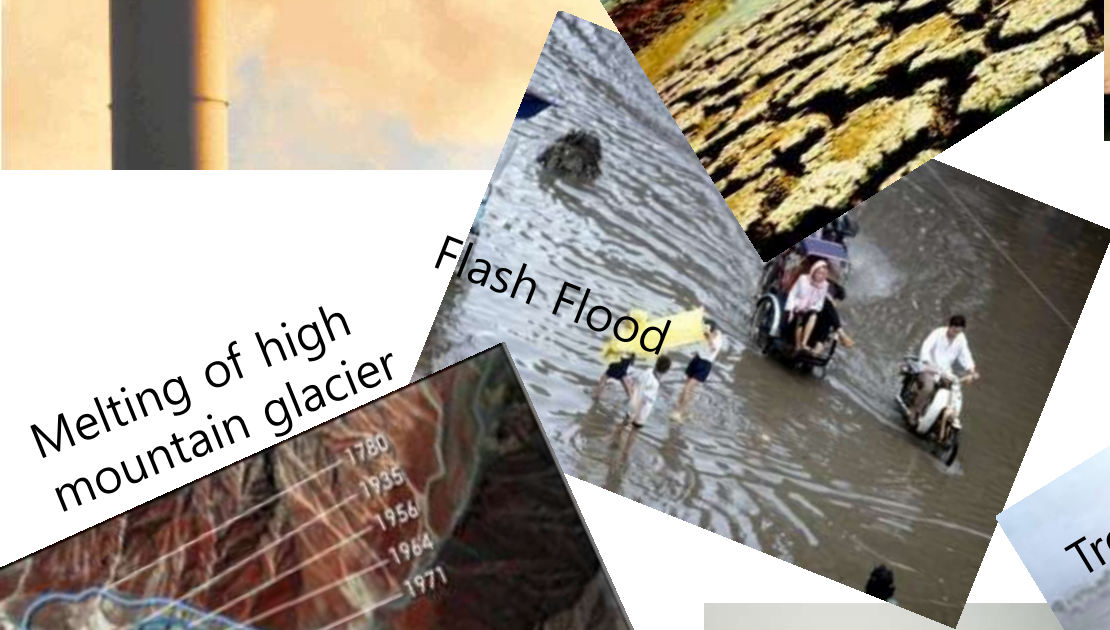
Industrialization



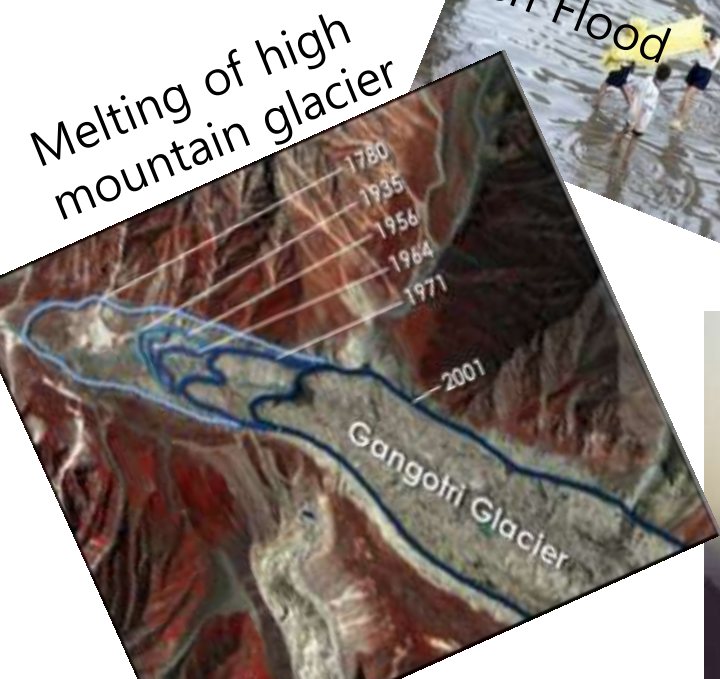
Drought



Dust storm



Flash Flood



Melting of high mountain glacier



Tropical cyclone



smok

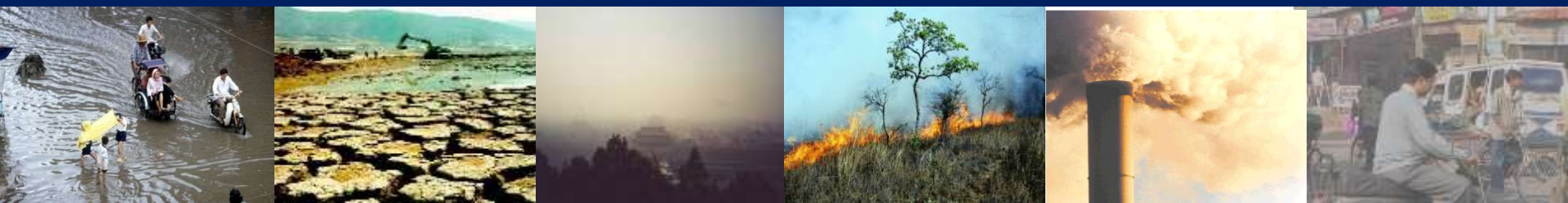


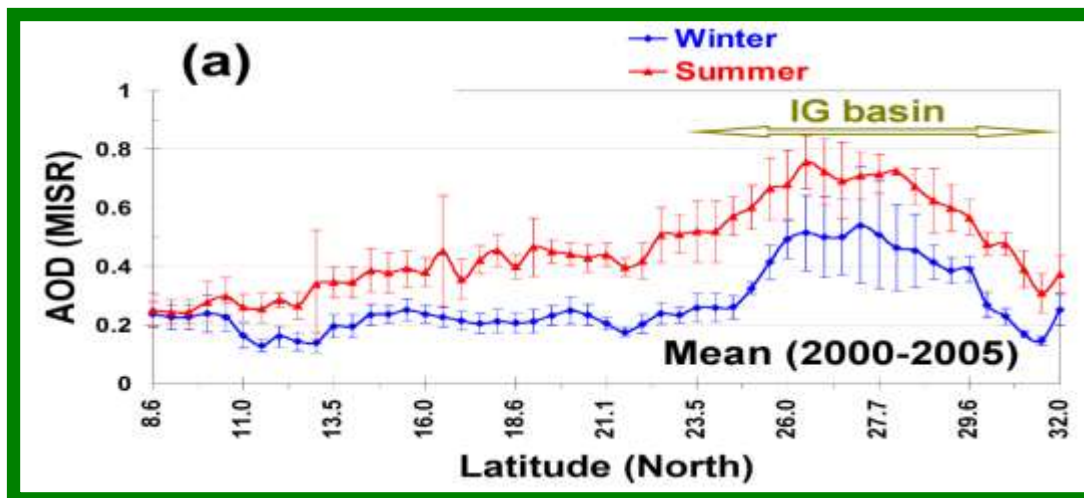
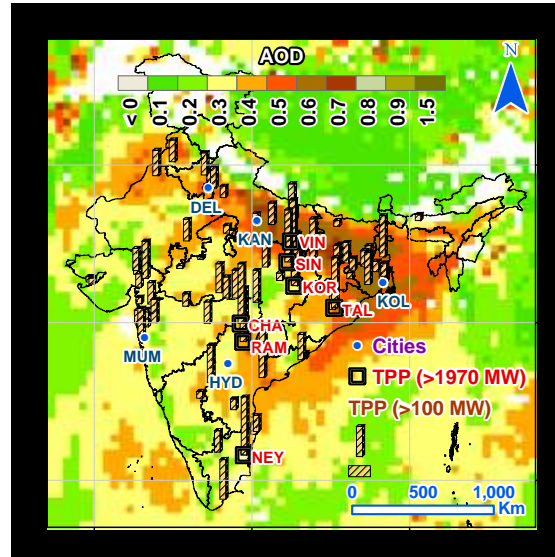
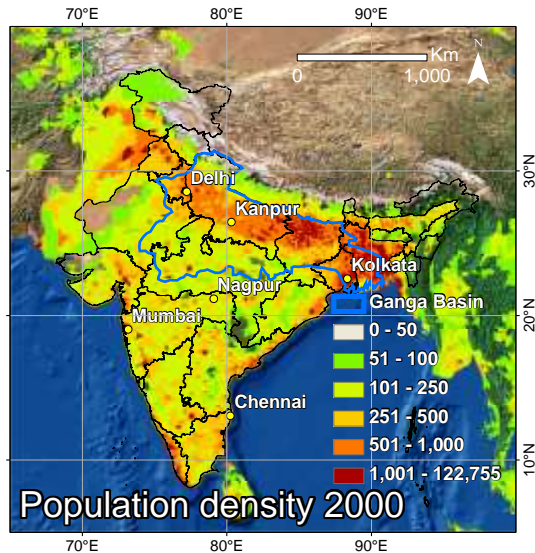
Biomass burning



# Aerosols and the Asian Monsoon

- Over 60% of world population live in the Asian monsoon regions
- **Monsoon** related droughts and floods, and **aerosols** are the two most serious environmental hazards in Asian monsoon regions,
- Sea surface temperature, and land surface processes drive the large-scale monsoon, through surface heating gradients and atmospheric heat sources and sinks.
- The monsoon water cycle is driven by atmospheric heating/cooling, through the dynamical interaction of wind, moisture, clouds and rainfall and surface processes.
- **Suspended particles (aerosol, clouds, precipitation) in the atmosphere regulate with the land-sea thermal gradient, interact with heat sources and sinks, and alter the monsoon water cycle**

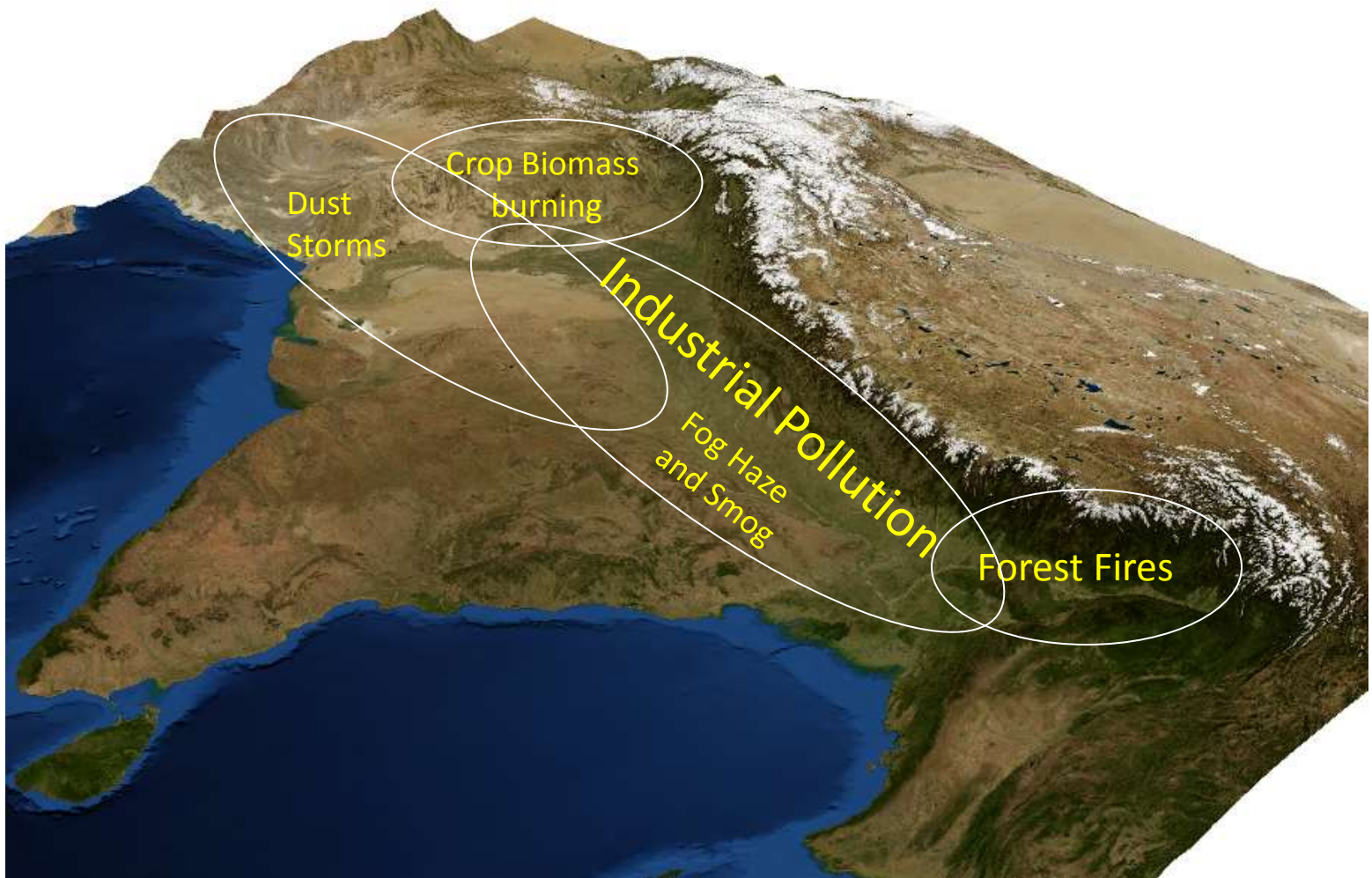




Courtesy  
Ramesh and Prasad

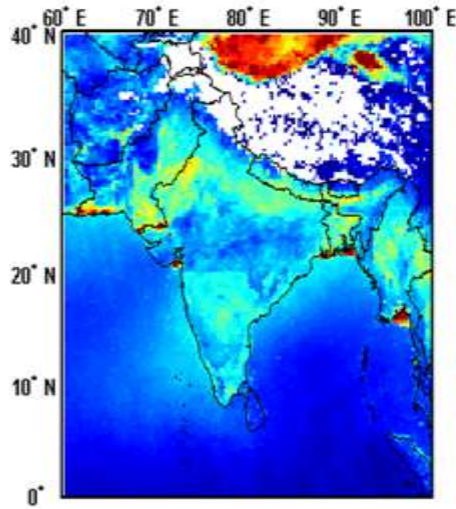


# Complex mixture of aerosols over the Indo-Gangetic Plain

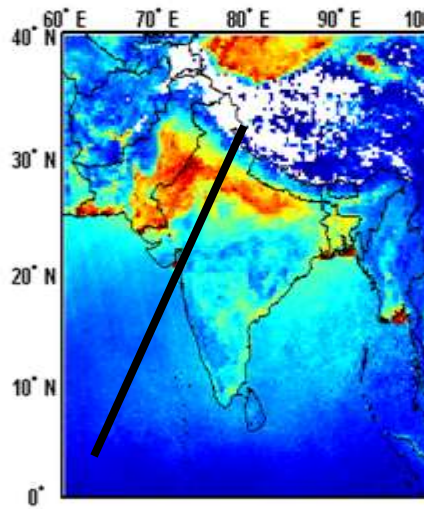




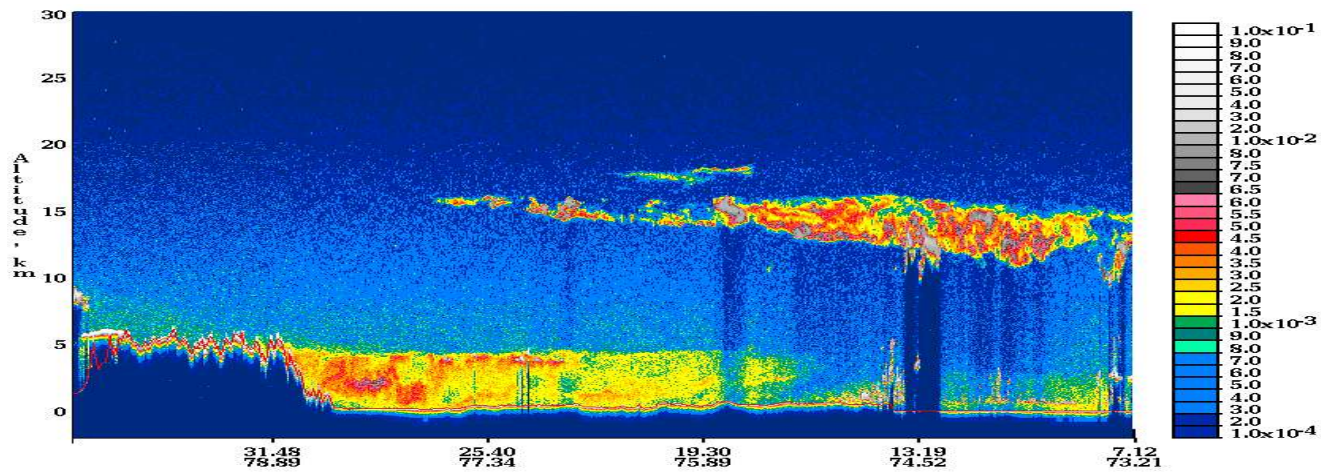
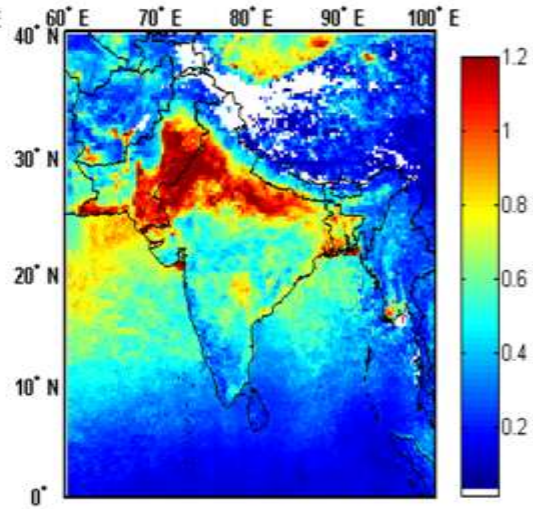
April



May

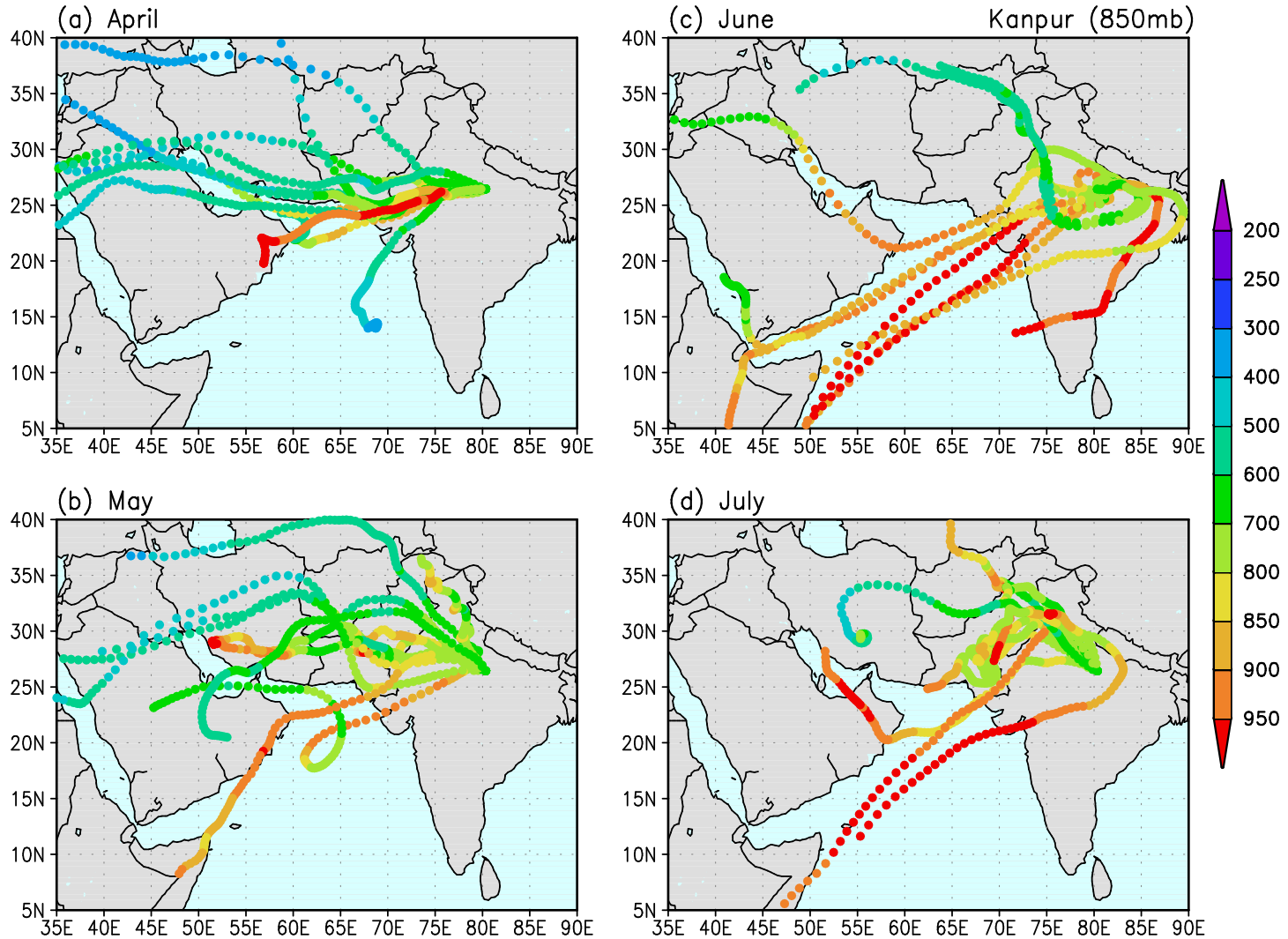


June





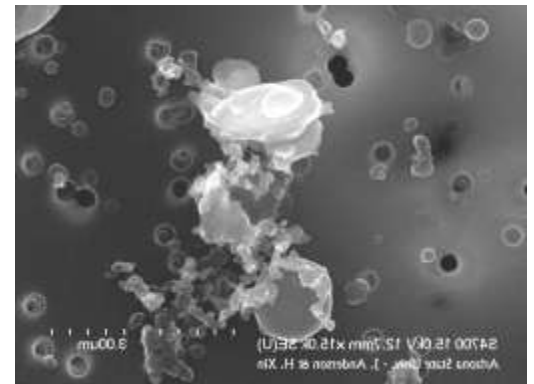
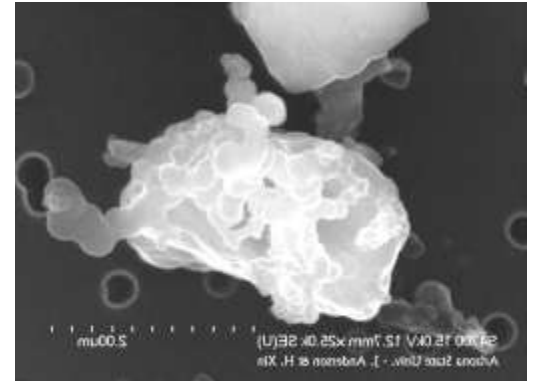
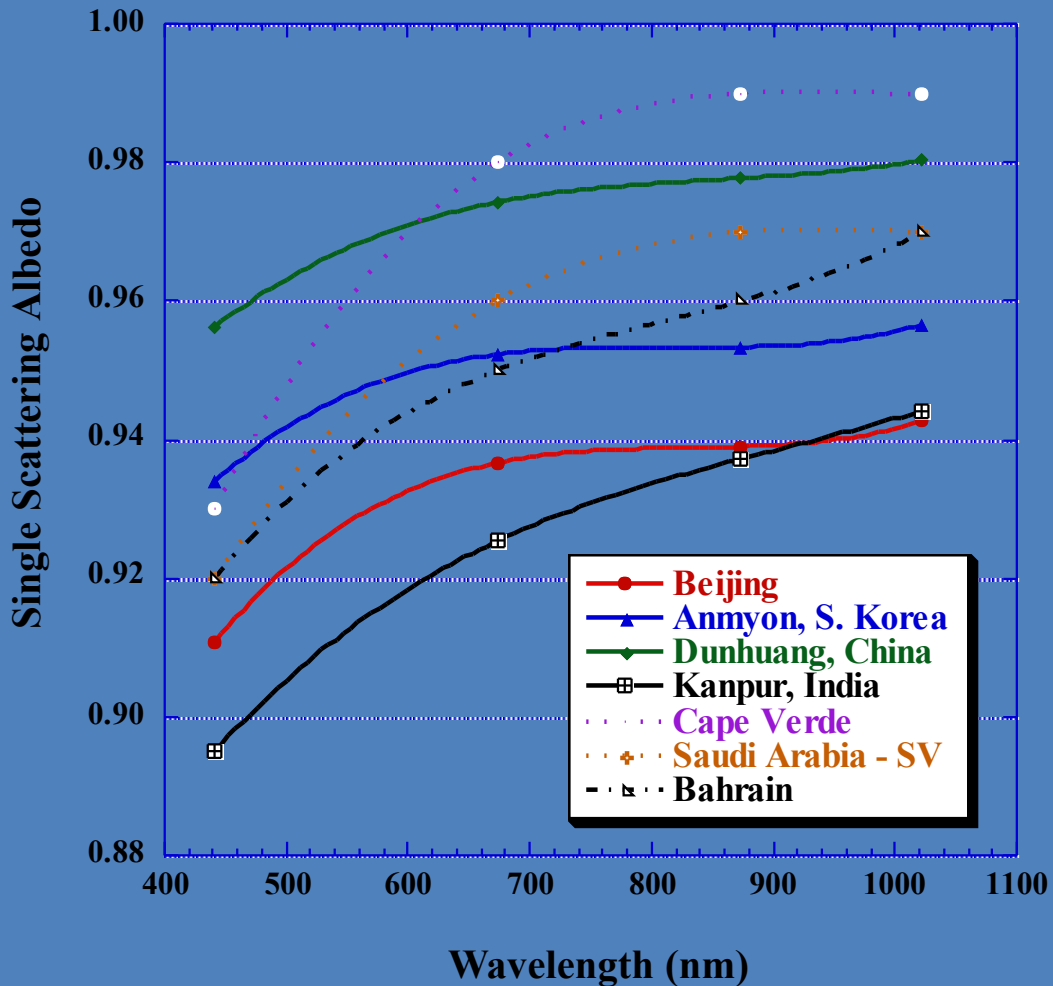
## 11-day Back Trajectories for 850hPa airmass at Kanpur, India , 2008





# How absorbing are dust aerosols ?

Comparison of Asian Dust Single Scattering Albedo versus Saharan and Saudi Arabian Dust



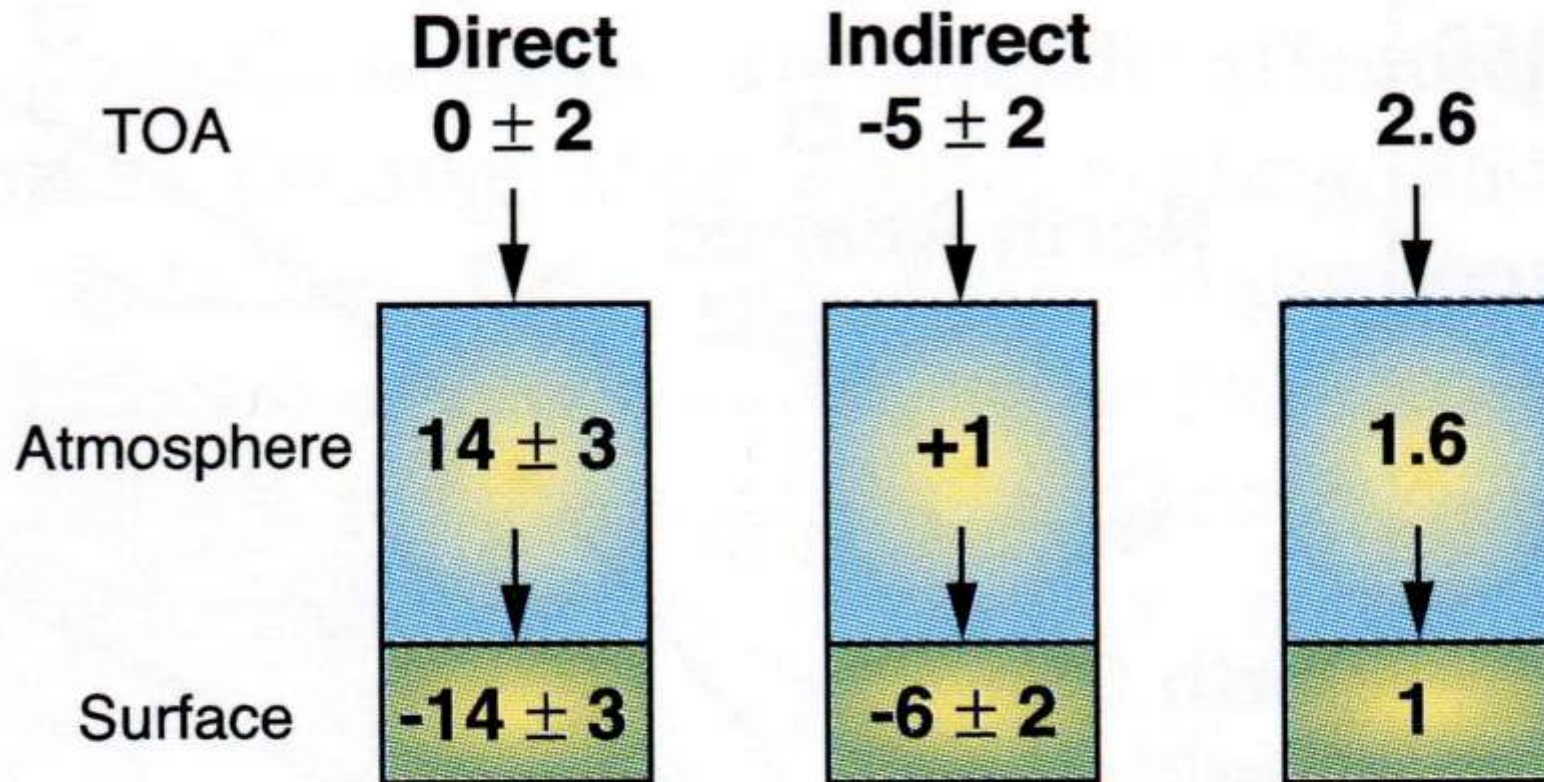
Micrographs of soot coating and aggregate on mineral dust from Asian dust (courtesy of J. Anderson)

**Dusts over major industrial cities are more absorbing (solar radiation)**

*ABCs and GHGs: Impact on Regional Radiation Budget*

**Aerosol  
forcing ( $\text{W m}^{-2}$ )**

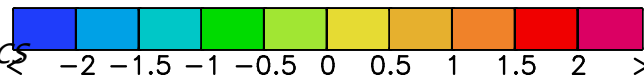
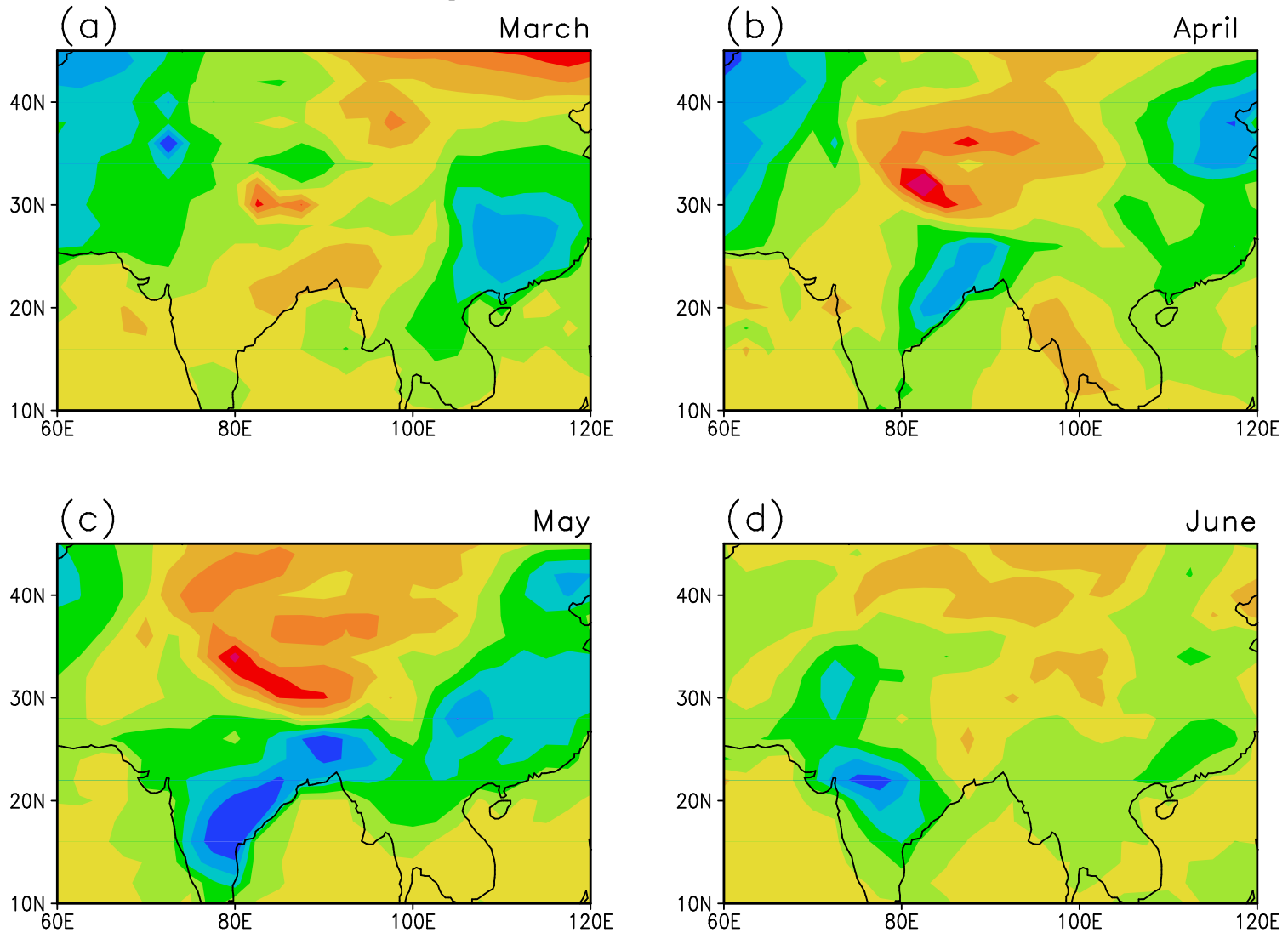
**Greenhouse  
forcing ( $\text{W m}^{-2}$ )**



**Tropical Indian Ocean: INDOEX  
(Preindustrial to 1996-1999; January to April)**



## Aerosol-induced anomalies of surface air temperature in NASA fvGCM

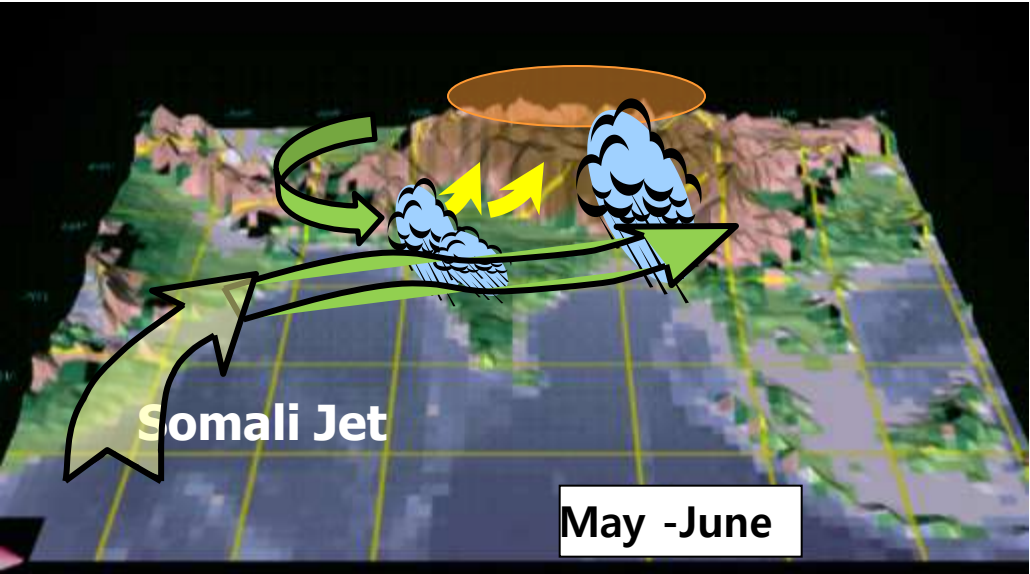


°C

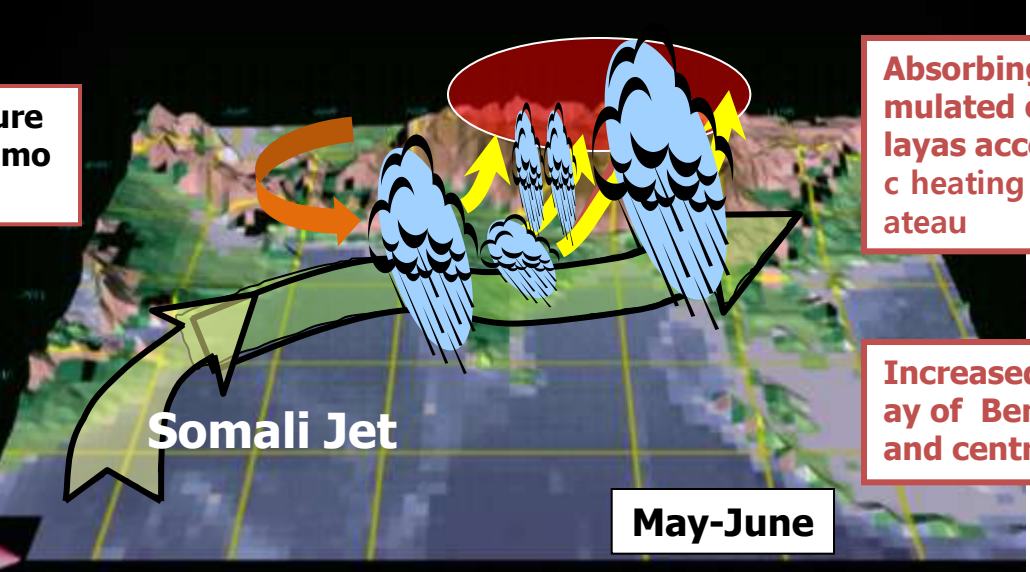


# The Elevated Heat Pump Hypothesis

(Lau et al. 2006, Lau and Kim 2006, Lau et al. 2008)



**Increase dust and moisture transport from low level monsoon westerlies**



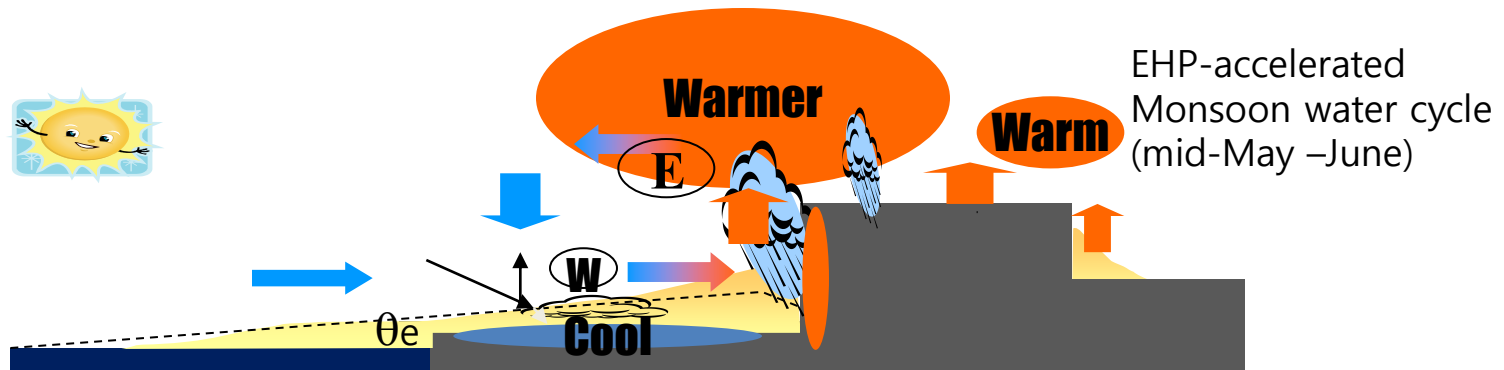
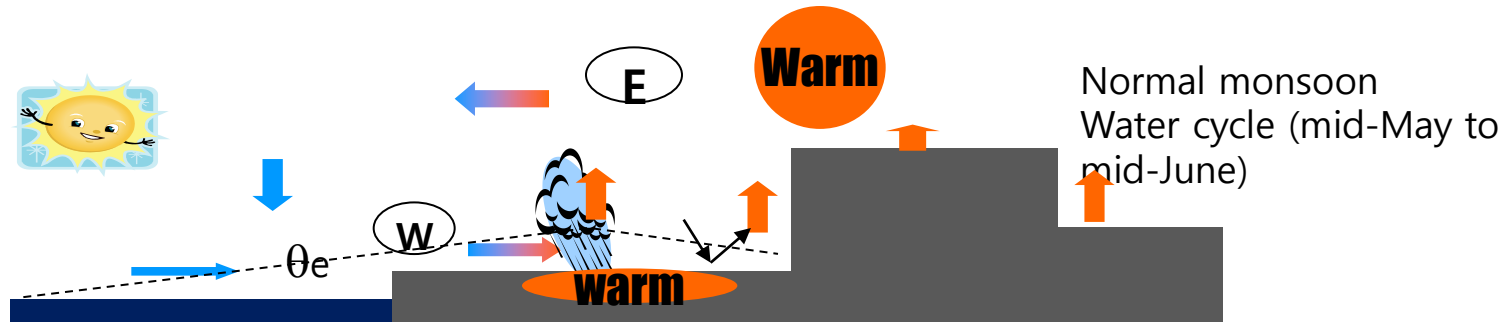
**Absorbing dust and BC accumulated on foothills of Himalayas accentuate atmospheric heating over the Tibetan Plateau**

**Increased convection over Bay of Bengal, eastern TP, NE and central India**



# The Elevated Heat Pump (EHP) hypothesis

(Lau et al. 2006, Lau and Kim 2006, Lau et al 2008)



## EHP postulates:

- Warming and moistening of the upper troposphere over the Tibetan Plateau
- An advance of the rainy season in northern India/Napal region in May-June
- The increased convection spreads from the foothills of the Himalayas to central India, resulting in an intensification of the Indian monsoon. in June-July
- Subsequent reduction of monsoon rain in central India in July-August
- Enhanced snowmelt and rapid retreat of mountain glacier



# Climate Change is a "clear and present" danger in high mountain regions:

Accelerated melting of Himalaya-Tibetan Plateau glacier will endanger the water supply of the entire Asia

The Himalaya-Tibetan Plateau Glaciers feed seven of Asia's great rivers: the Ganga, Indus, Bramaputra, Salween, Mekong, Yangtze, and Yellow River that supply fresh water for over 60% of the world population





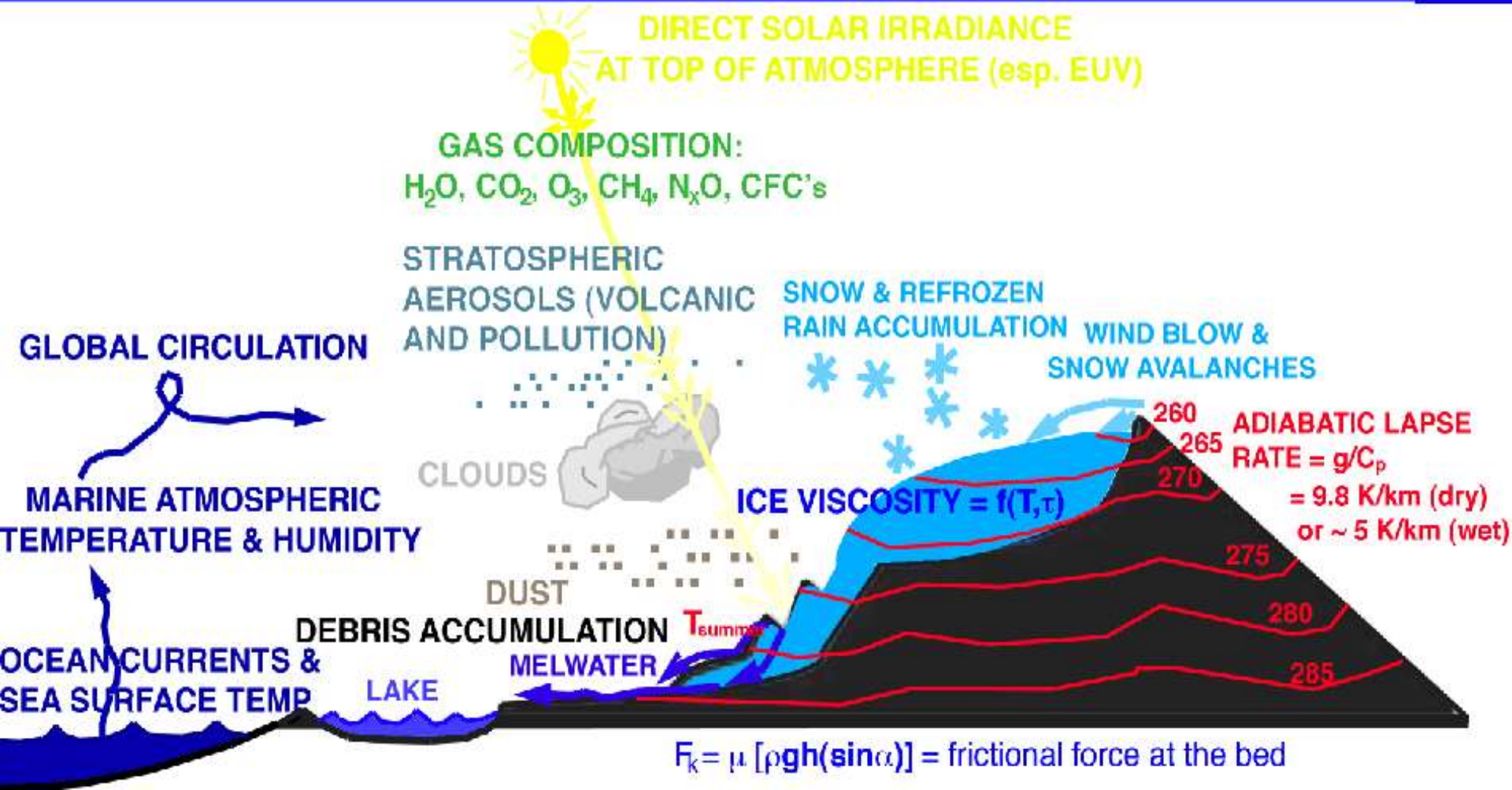
## Accelerated melting of Himalaya-Tibetan Plateau glacier will endanger the water supply of the entire Asia!

The Himalaya-Tibetan Plateau Glaciers feed seven of Asia's great rivers: the Ganga, Indus, Bramaputra, Salween, Mekong, Yangtze, and Yellow River that supply fresh water to many countries in Asia.

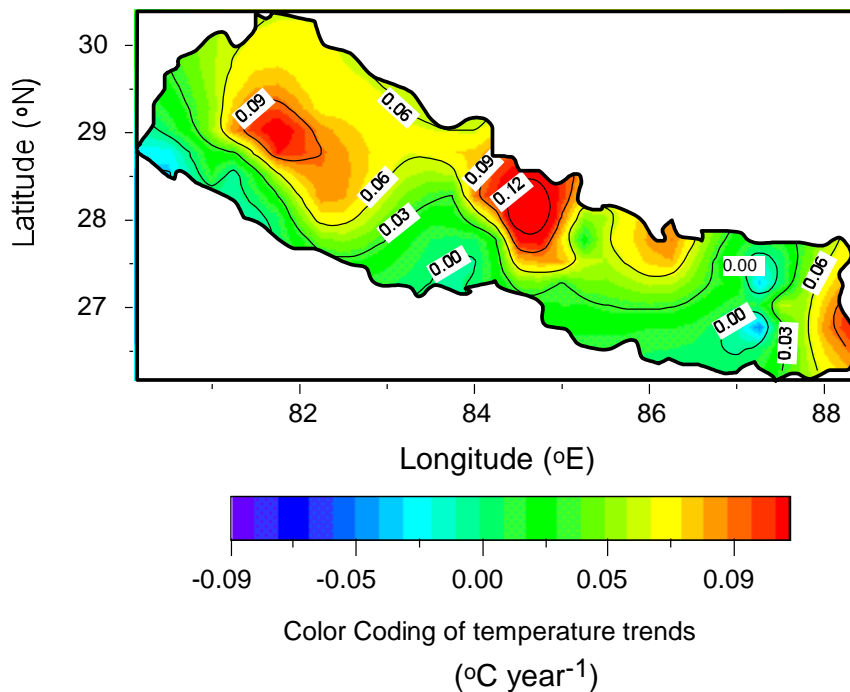




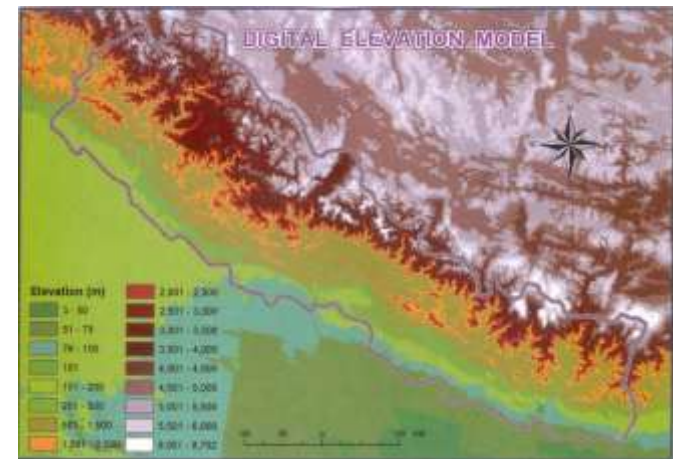
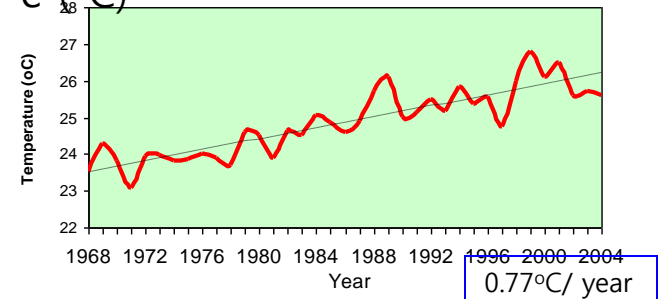
# Climate and environmental change related to glaciers is more than about rising temperatures



# Spatial Distribution of maximum temperature trends in Nepal for the period 1977 to 1994.



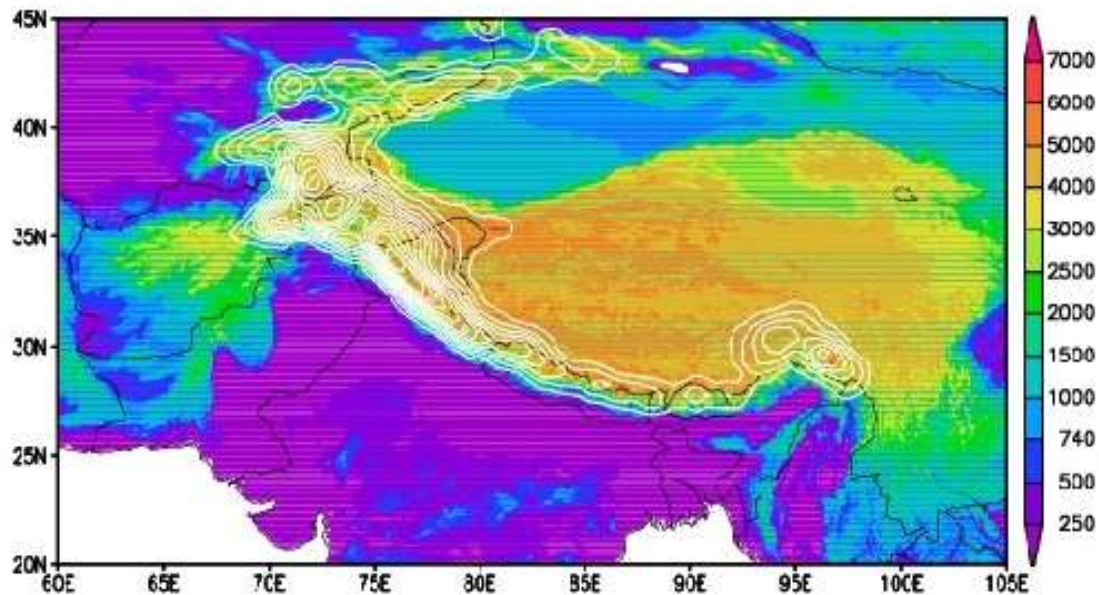
## Kathmandu Maximum Temperature (°C)



Courtesy of Dr. Madan Shrestha, DG, Met. Service, Nepal



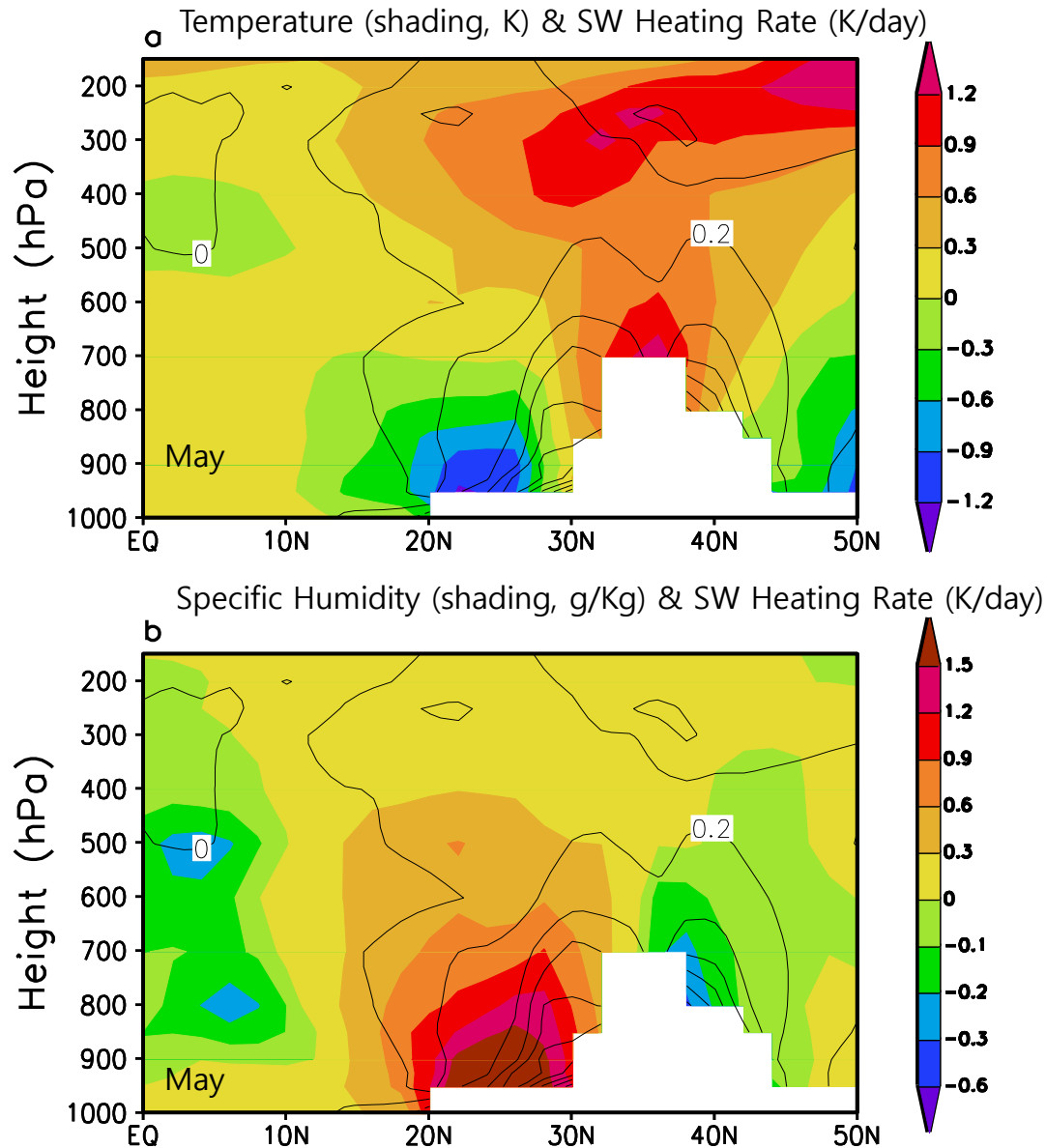
# Possible causes of accelerated melting of snowpack in the Himalayas and Tibetan Plateau



- Greenhouse Warming
- Atmospheric heating due to aerosols (EHP)
- Snow-darkening effect from soot and dust, and other impurities

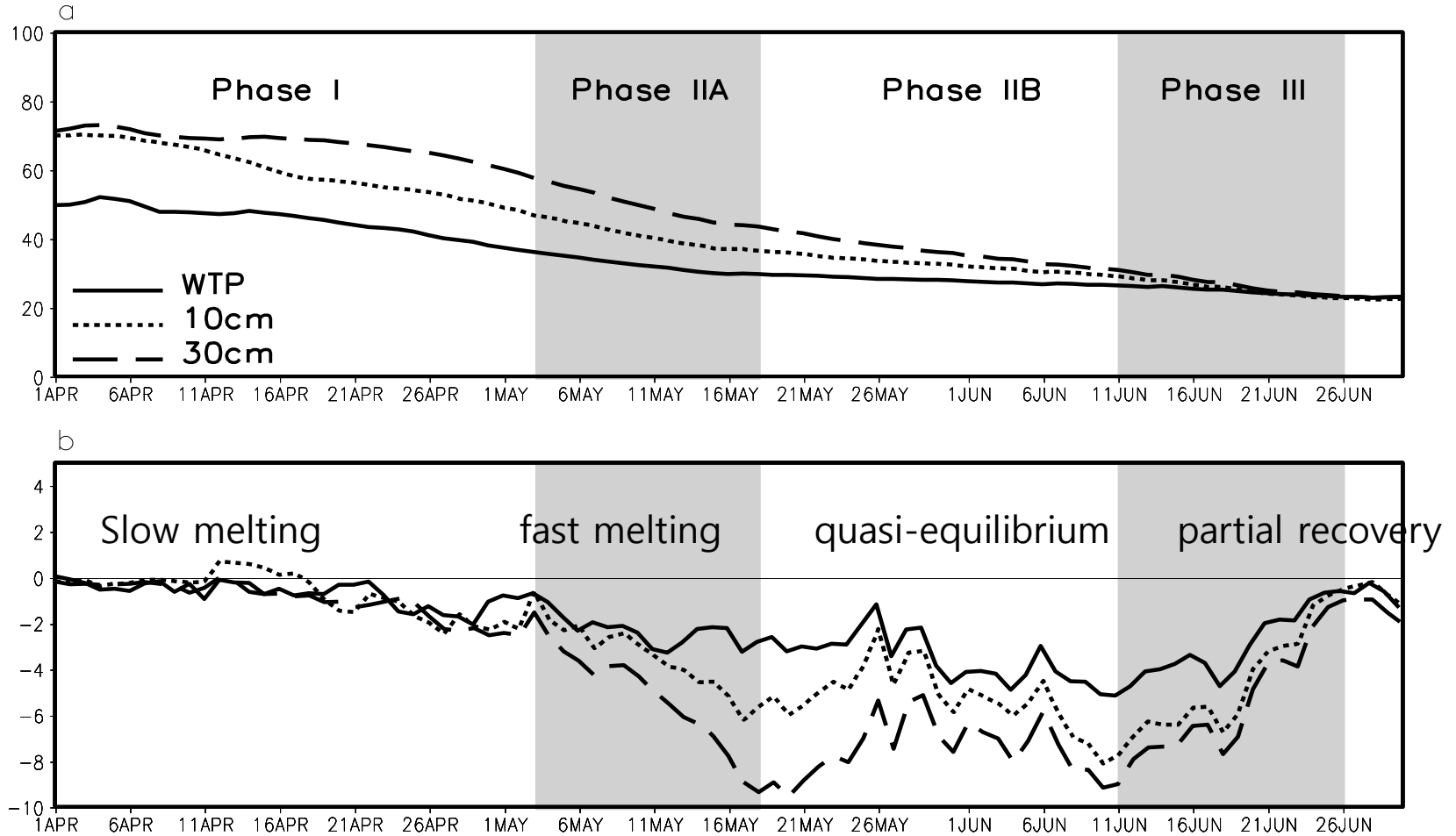


## EHP induced heating and moistening along 80°E





Relative percentage change of surface albedo (due to changes  
in snow-cover, and exposure of underlying land surface )  
from NASA climate model





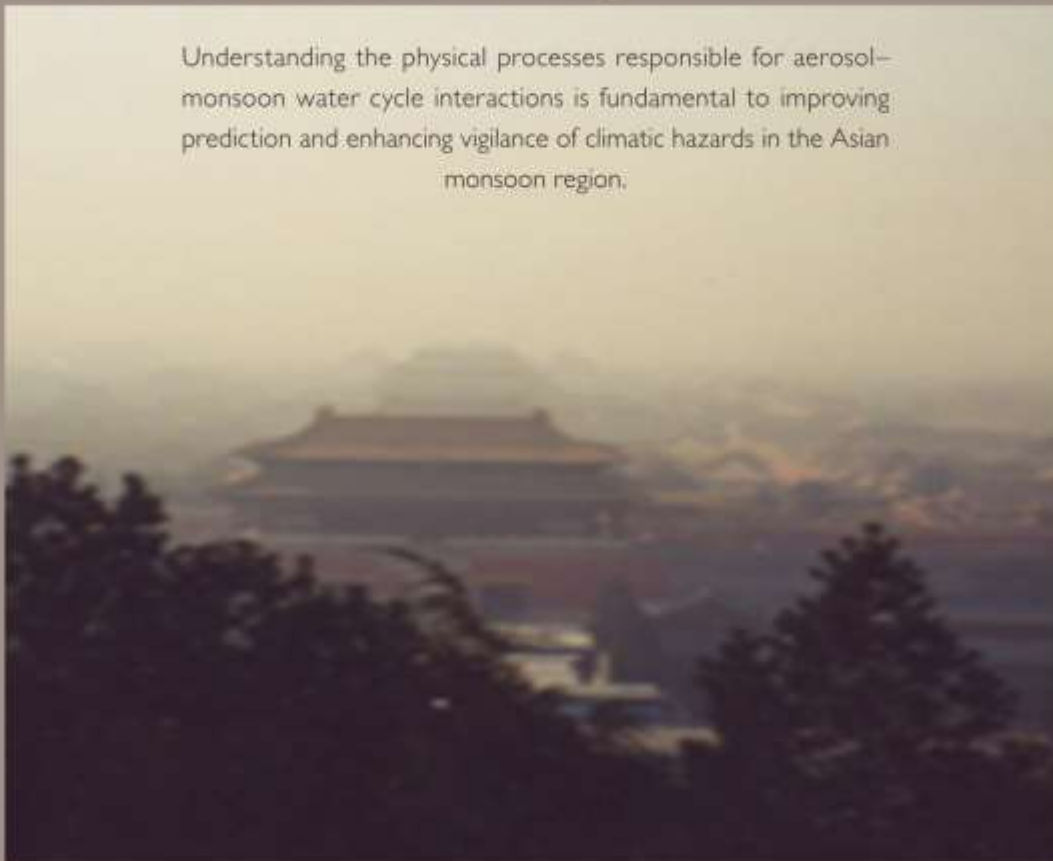
*Bull. Am. Meteor. Soc., 2008*

# THE JOINT AEROSOL– MONSOON EXPERIMENT

A New Challenge for Monsoon Climate Research

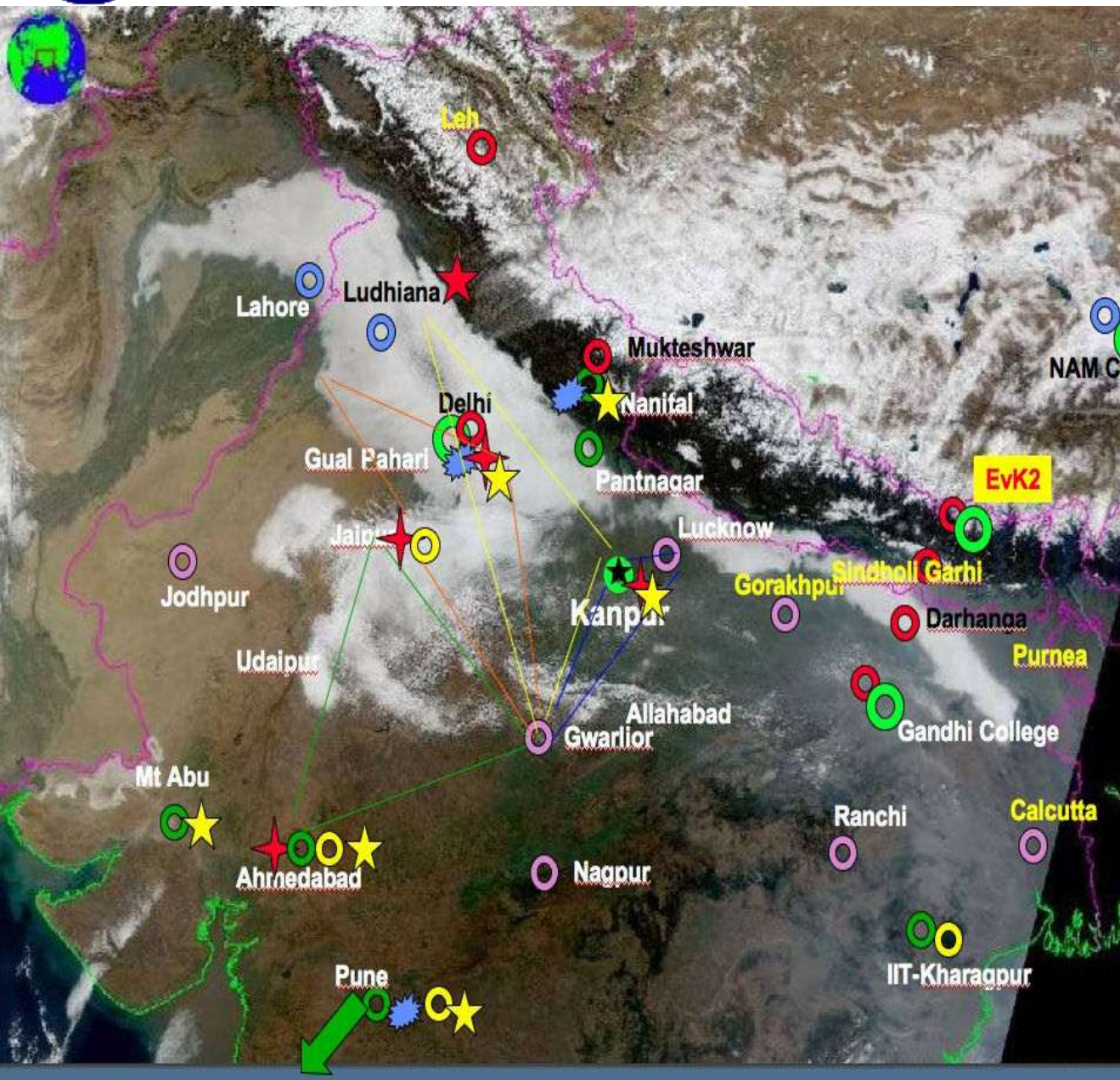
BY K.-M. LAU, V. RAMANATHAN, G.-X. WU, Z. LI, S. C. TSAY, C. HSU, R. SIKKA, B. HOLBEN, D. LU,  
G. TARTARI, M. CHIN, P. KOUDELOVA, H. CHEN, Y. MA, J. HUANG, K. TANIGUCHI, AND R. ZHANG

Understanding the physical processes responsible for aerosol–monsoon water cycle interactions is fundamental to improving prediction and enhancing vigilance of climatic hazards in the Asian monsoon region.



Smoke consisting of mixtures of dust and industrial pollution covering the Forbidden City, Beijing, China.

Air pollution and monsoon floods and droughts are have found that anthropogenic aerosols may signifi-



## JAMEX/TIGERZ - 2009

★ SMART-COMMIT: Proposed 2009

○ AERONET: Operational

○ Microtop: Planned



Lidar



Chemistry

○ Microtops: Operational

○ AERONET: Potential

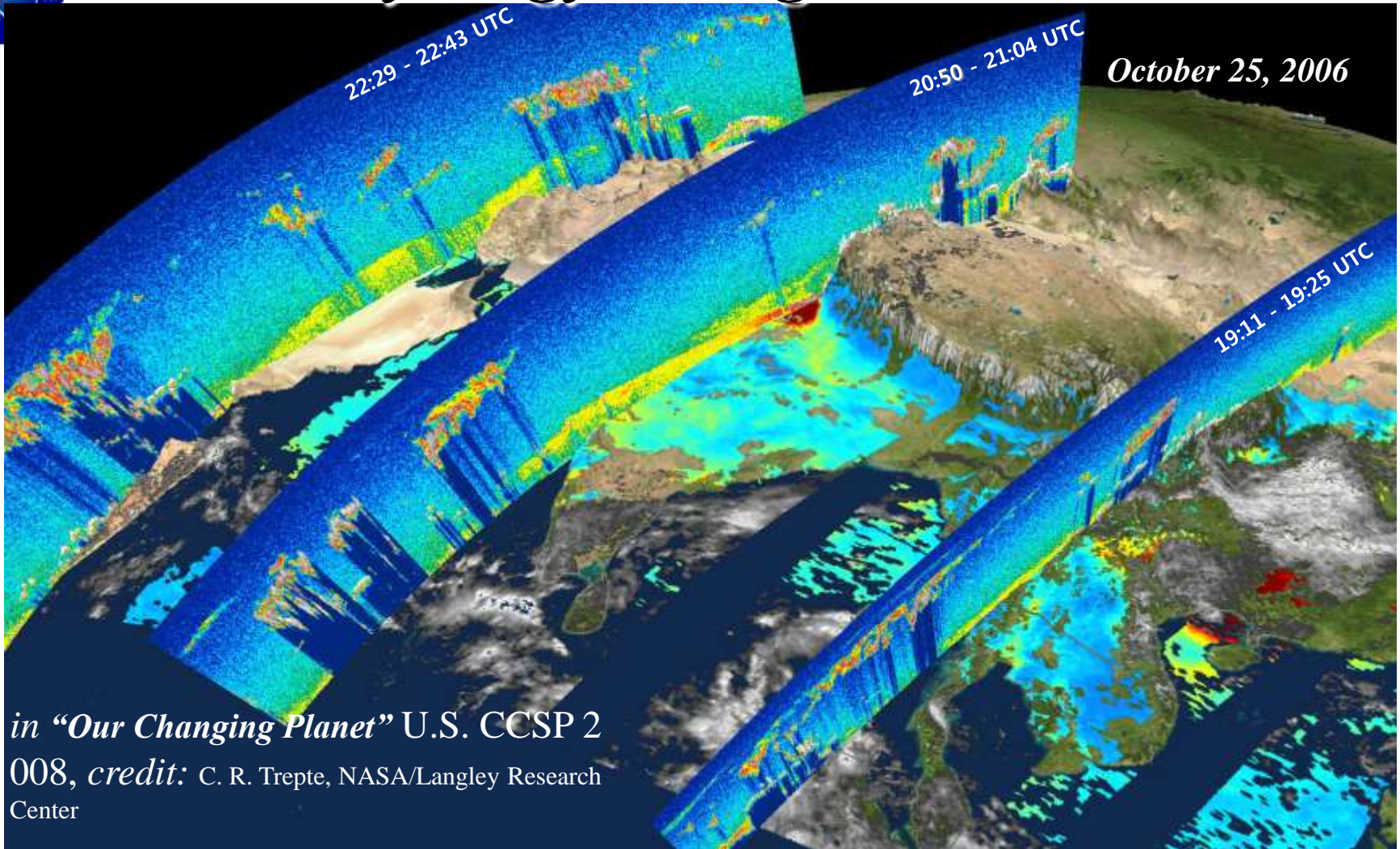
○ Radiosonde

★ Solar Flux

★ MWR



# Synergy: Imager + Lidar



- Lau *et al.*, 2008, *Bull. Amer. Meteor. Soc.*, 8, 9, doi:10.1175/BAMS-89-3-369.
- A *three-dimensional* perspective of aerosol/cloud distributions
- Enhance study of *Elevated-Heat-Pump* and *aerosol-cloud-glacier* interaction



The highest aerosol observatory, NCO-P (ABC-Pyramid):  
the southern slope of the Himalayas (5079 m a.s.l.)

Meteorological parameters, Optical Particle Counter (**OPC; 0.25-32  $\mu\text{m}$** ), Scanning Mobility Particle Sizer (**SMPS; 10.31-669.8 nm**), and **equivalent Black Carbon concentration (MAAP)** data in 2006 at NCO-P site were observed (Bonasoni et al., 2008) and used in this study.



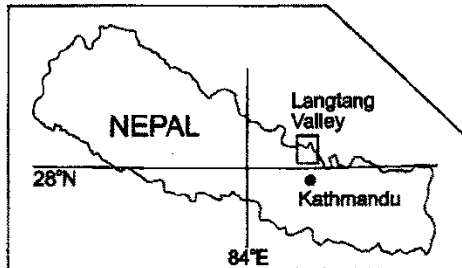
NCO-P site (ABC-Pyramid site): <http://evk2.isac.cnr.it/>



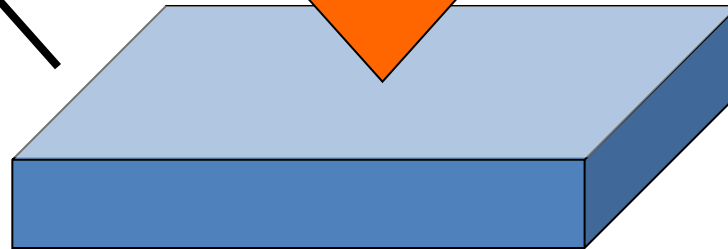
# Estimated “lower limit” of effect of BC deposition in darkening Himalayan glaciers

BC deposition of **209**  $\mu\text{g}/\text{m}^2$  by dry fallout  
during March-May in 2006 ( $2.27 \mu\text{g m}^{-2} \text{ day}^{-1}$ )

Surface snow density  
at Yala glacier: **195 to 512 kg/m<sup>3</sup>**  
(Fujita et al., 1998)



2cm surface snow



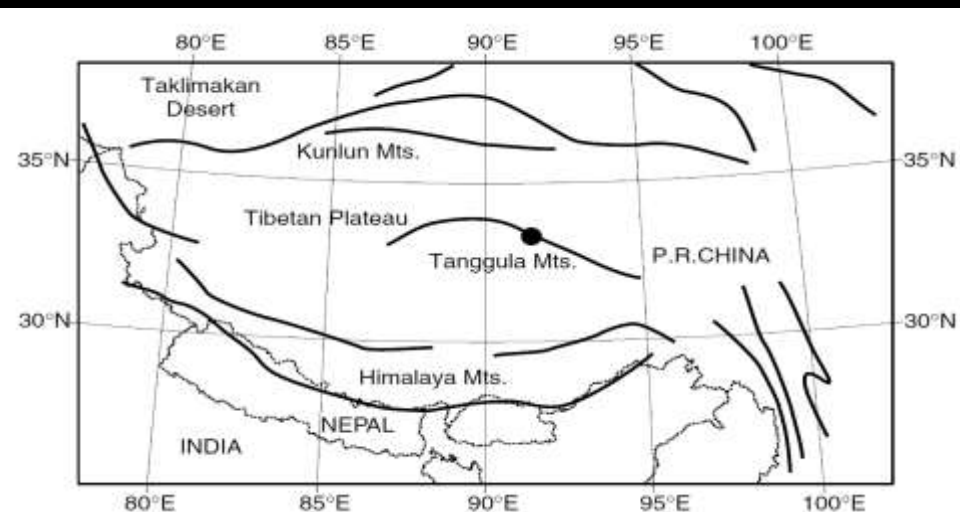
$$v_s = \frac{2}{9} \frac{r^2 (\rho_{BC} - \rho_{Air}) g}{\eta}$$

Now, imagine that we take snow samples from 2cm surface snow layer. At this time, we assume slight BC layer on the pure ice with no contamination.

# Estimation of additional annual runoff from snow-darkening effect



Figure 3. Photograph of the Da (left) and Xiao (right) Dongkemadi Glaciers



- Using a glacier mass balance model (Fujita, 2007, Fujita et al., 2007). with the same input data as used in previous studies
- Over a typical Tibetan glacier (Dongkemadi glacier), continuous albedo reductions of 1.6 & 4.1% may increase by 9.0-24.8% the annual run off from discharge of melted snow

- ◆ More Westerly dominated precipitation
- ◆ Less monsoon-influenced precipitation
- ◆ Glaciers grow by winter accumulation
- ◆ Less glacier disintegration & lake growth
- ◆ EHP net influence small?\*

- ◆ Less intense melting, more intense sublimation
- ◆ More cold-based ice
- ◆ Less debris cover
- ◆ Spatial variability of Elevated Heat-Pump effect
- ◆ Less soot-affected exposed ice surfaces
- ◆ But more exposed ice to be affected
- ◆ More sensitive to precipitation changes and wind

- ◆ More intense melting
- ◆ More warm-based ice
- ◆ More debris cover
- ◆ Strong Elevated Heat Pump effect
- ◆ More soot effect on exposed ice surfaces
- ◆ But less exposed ice to be affected
- ◆ Glaciers are more sensitive to warming

- ◆ Less Westerly dominated precipitation
- ◆ More monsoon dominated precipitation
- ◆ Grow mainly by summer snow accumulation
- ◆ More lake growth and glacier disintegration
- ◆ Strong EHP reduces glacier stability\*

Contact: Jeff Kargel

MODIS base image courtesy of GSFC/NASA

• Glacier behavior varies across the region, with faster retreat in the east. Possibly glaciers in northwest pick up more snow precipitation due to Elevated Heat Pump (EHP) and other climate mechanisms thus partly offsetting heating/melting. Glaciers in the eastern Himalaya may be more sensitive to EHP heating and are melting more quickly.



## *Thoughts on Mitigation/Adaptation Strategy*

- 1)  $\text{CO}_2$  has long-life time of centuries or longer; effects are long-term  
Doing nothing now, will cost us more later.
- 2) The observed global rise in temperature is likely a delayed response to industrialization years past; even if we stabilize  $\text{CO}_2$ , warming will continue for more than a century.
- 3) Regionally, climate change science is still unfolding. Asia is potentially the most vulnerable region: aerosol, monsoon floods/drought, high-mountain glaciers; fresh water supply ....
- 4) In Asia, availability of fresh water is a "clear and present danger" and is likely to get worse if nothing is done. Mitigation or adaptation strategy should be focused on alleviating hazards and potential conflicts arising from fresh water shortage
- 5) Politically, it would be wise to take actions that can produce benefits in the near term, i.e., reduce BC emission, and deal with short-lived GHG's first, e.g.  $\text{CH}_4$ ,  $\text{O}_3$ ,  $\text{NO}_x$ .
- 6) Geo Engineering, e.g. injecting sulphate in lower stratosphere.....  
Don't go there. *Beware of the "Law of Unintended Consequences" !*



# NASA climate model simulation of accumulation of soot over the Indo-Gangetic Plain

Aug 01, 2009  
00:00 UT

Black Carbon Aerosol Optical Thickness



0.002 0.005 0.01 0.02 0.05 0.1 0.2

A horizontal color scale bar with a gradient from dark purple to light yellow. The values are marked at 0.002, 0.005, 0.01, 0.02, 0.05, 0.1, and 0.2.

[Video Link](#)



# Back Up

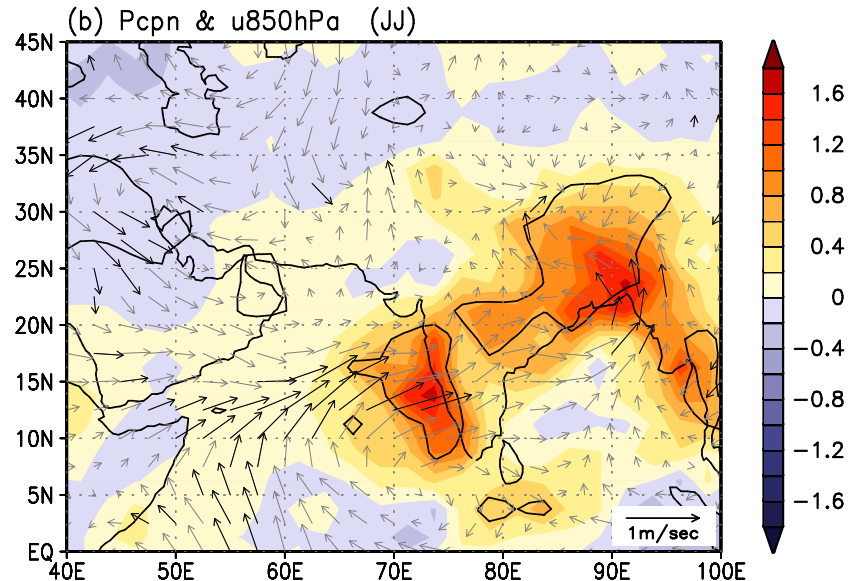
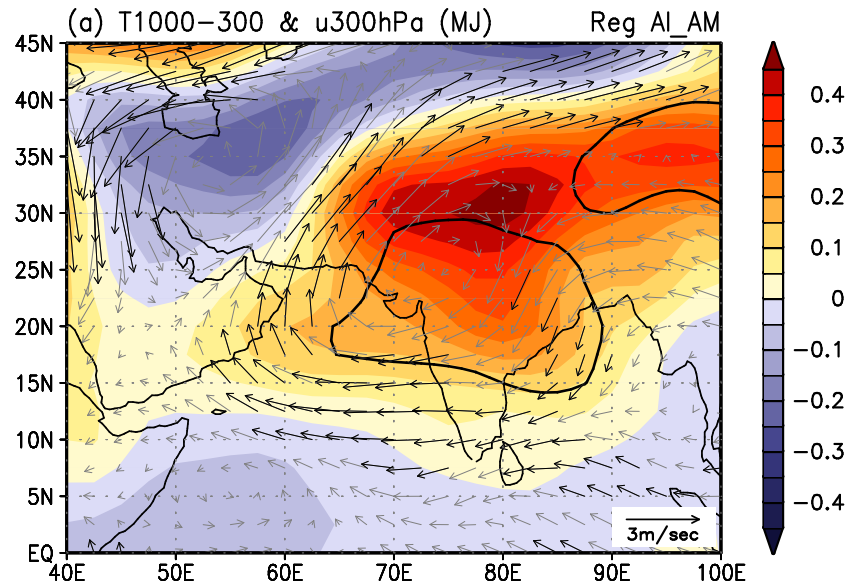


## Characteristic EHP large scale pattern from observations :

Enhanced tropospheric warming over the Tibetan Plateau and increased monsoon rainfall and winds in June-July, following major build-up of absorbing aerosol (dust and BC) over the Indo-Gangetic Plain in April-May

Data source: TOMS AI, GPCP, and NCEP re-analyses

Lau and Kim (2006, *GRL*)



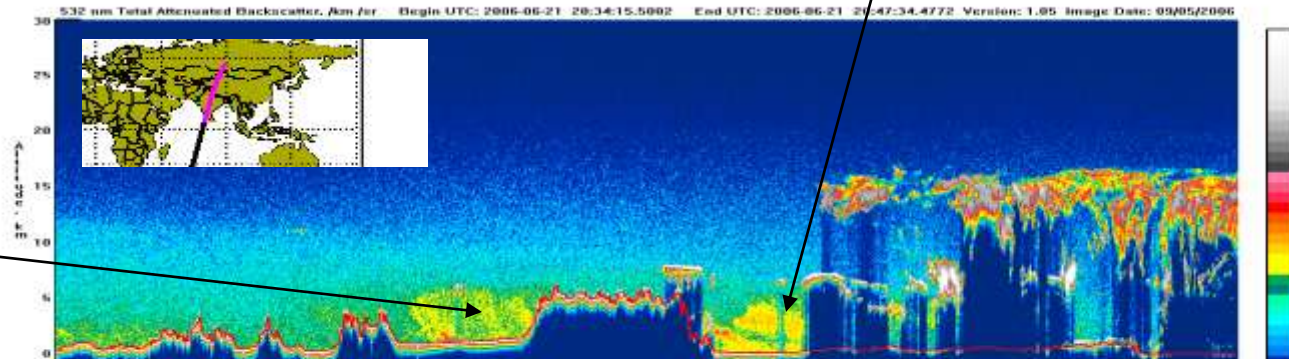


## Meridional cross-section of aerosol concentration from Calips

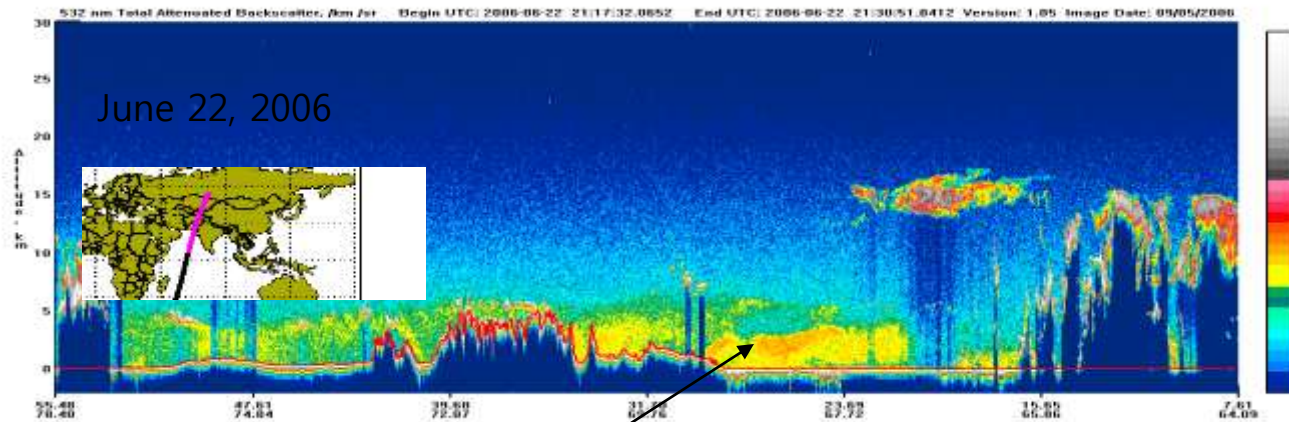
June 21, 2006

Indo-Gangetic Basin

Taklamakan Desert



June 22, 2006



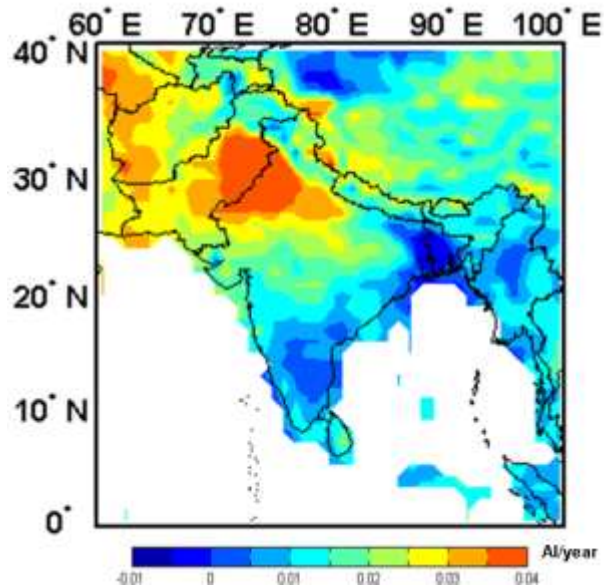
northeastern IGB

Reference"

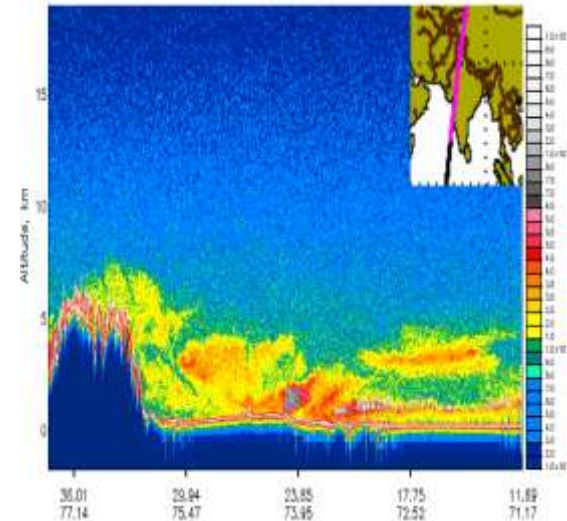
Lau, K.-M., K.-M. Kim (2008), Absorbing aerosols enhance Indian summer monsoon rainfall. iLEA PS Newsletter, No. 5, 22-24.



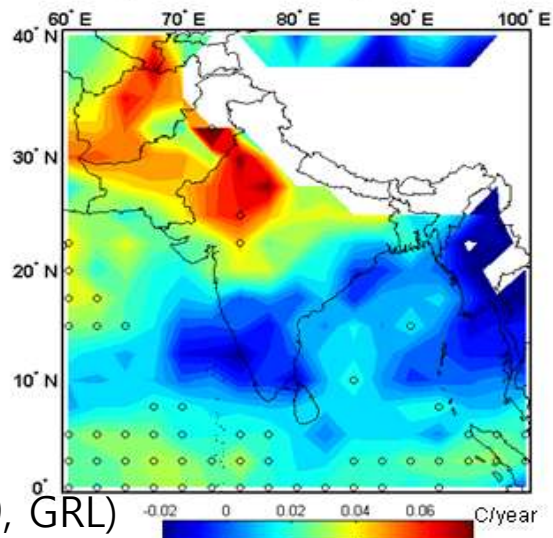
TOMS AI trend for May  
(1979-2001)



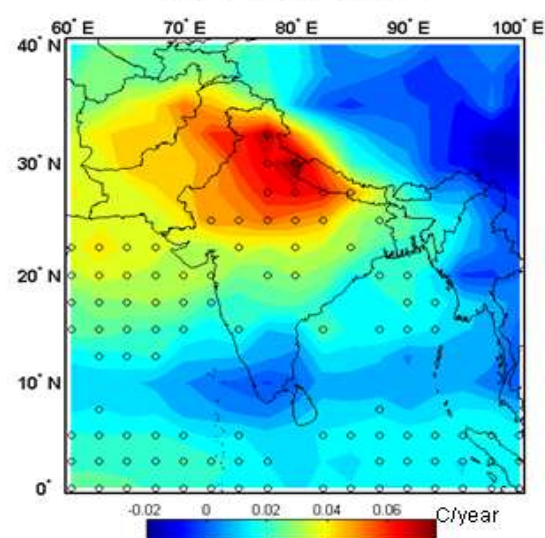
Aerosol profile from Calipso



Lower-Tropospheric Temperature  
(surface-4 km) Trend



Mid-Tropospheric Temperature  
(4-7 km) Trend





## Photo Gallery



### Tibetan Plateau

Glaciers that feed great Asian rivers are shrinking. See photographs by Jonas Bendtsen.

## Video



### Behind the Photo

Sometimes called the Third Pole, the region is a lockbox of snow and glacial ice that supplies fresh water to nearly a third of the world's people.

## Interactive

### Freshwater 101

We live on a planet covered by water, but how much is freshwater?

## Tibetan Plateau

## National Geographic

Published: April 2010



### The Big Melt

Glaciers in the high heart of Asia feed its greatest rivers, lifelines for two billion people. Now the ice and snow are diminishing.

By Brook Larmer

Photograph by Jonas Bendtsen

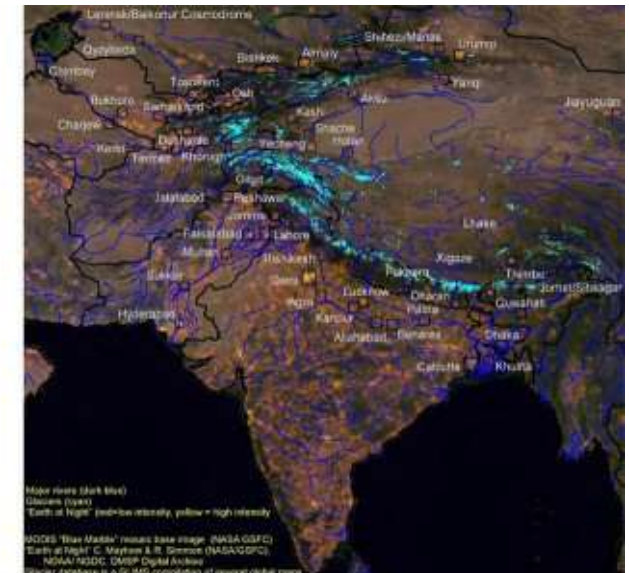
The gods must be furious.

It's the only explanation that makes sense to Jia Son, a Tibetan farmer surveying the catastrophe unfolding above his village in China's mountainous Yunnan Province.

"We've upset the natural order," the devout, 52-year-old Buddhist says. "And now the gods are punishing us."



- 





## *Afternoon Constellation of the “A-Train”*

1:38 PM

1:30 PM

1:15 PM



OMI - Cloud heights

OMI & HIRLDS - Aerosols

MLS& TES - H<sub>2</sub>O & temp profiles

MLS & HIRLDS - Cirrus clouds

Cloudsat

CALIPSO

MODIS/CERES - IR Properties

AIRS - Temperature and Humidity

Aqua

PARASOL

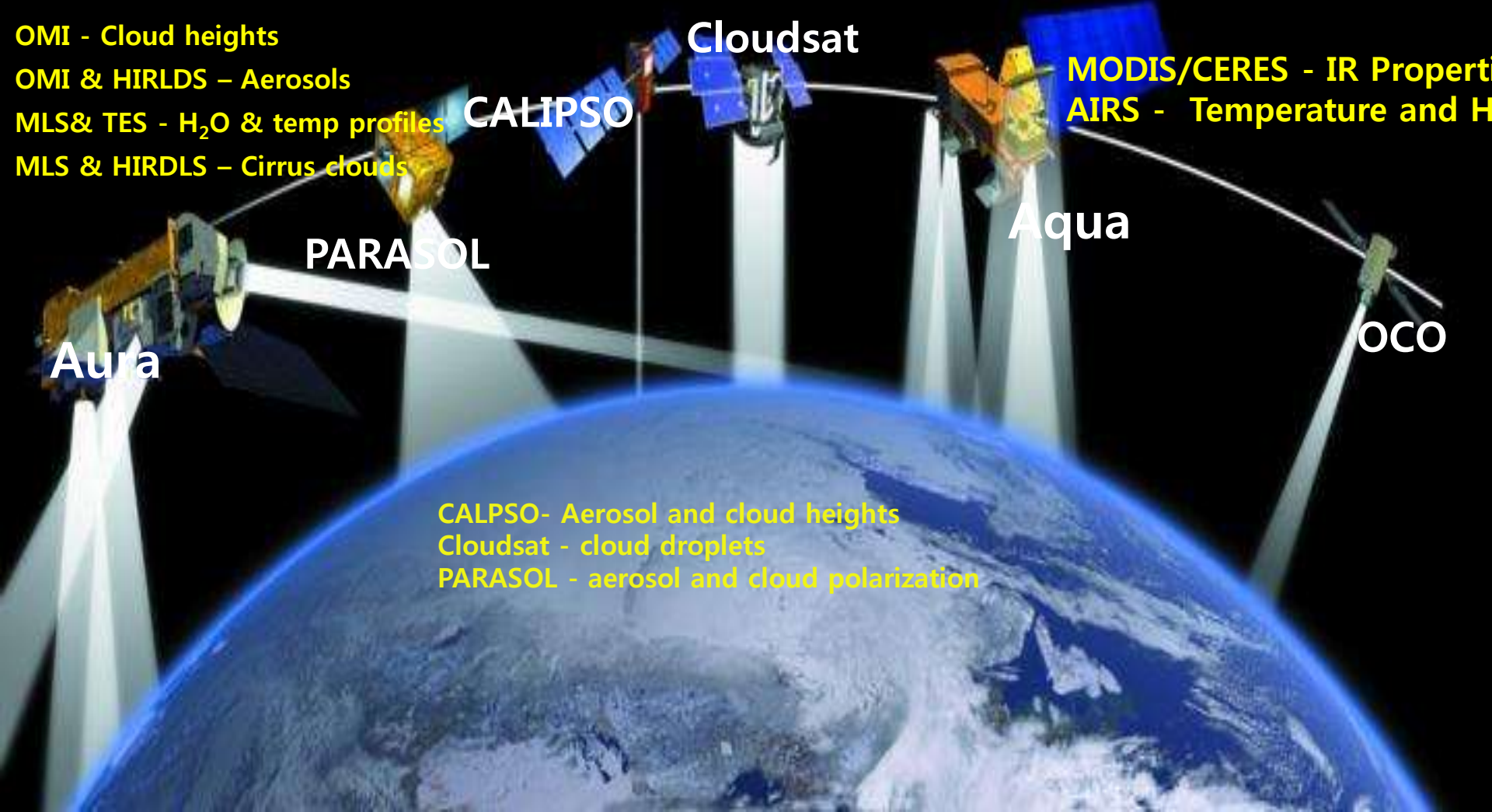
Aura

OCO

CALIPSO- Aerosol and cloud heights

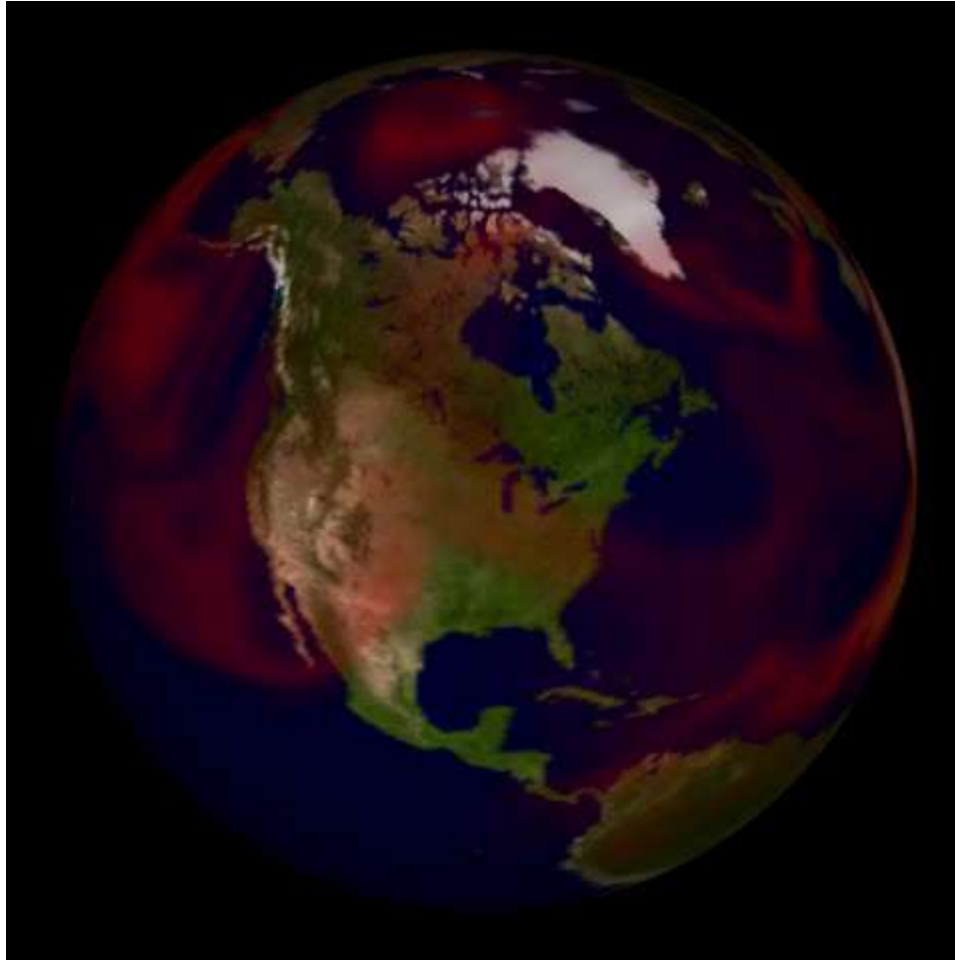
Cloudsat - cloud droplets

PARASOL - aerosol and cloud polarization

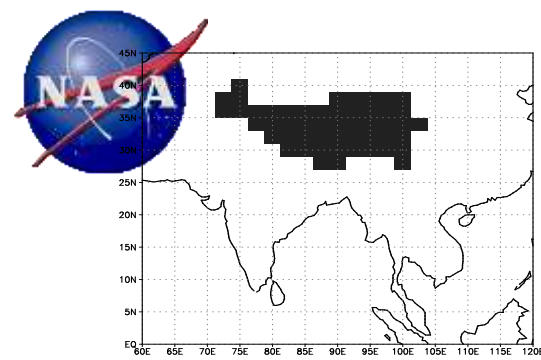




We need High-Resolution GCM or RCM with Interactive Aerosols with realistic aerosol microphysics to better understand Aerosol- Monsoon Water Cycle Interactions

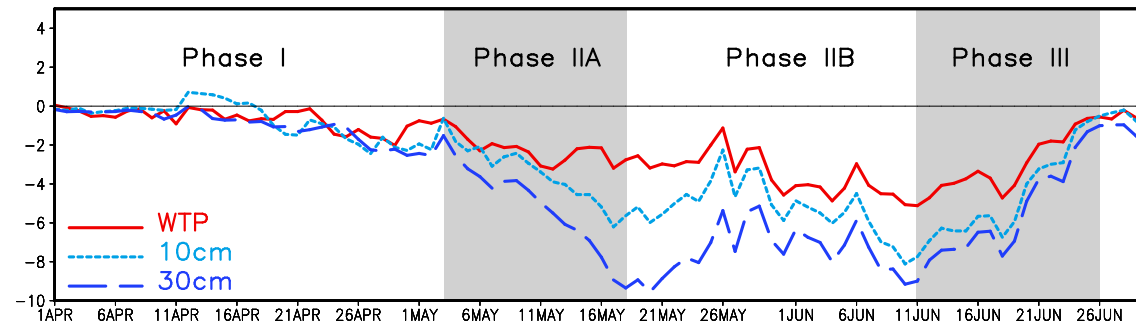


NASA GEOS5 model 5-day forecast of dust emission and transport (May 1-5, 2006)

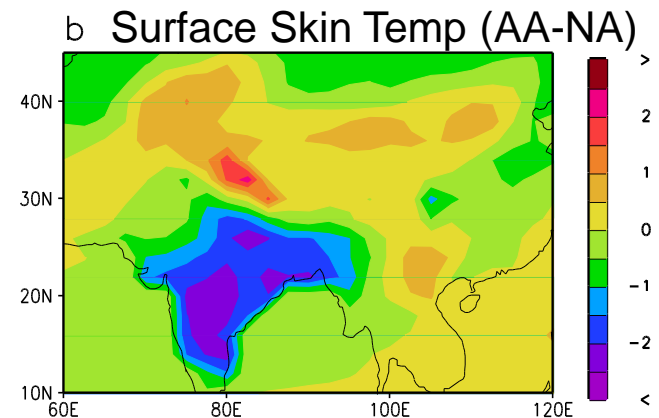
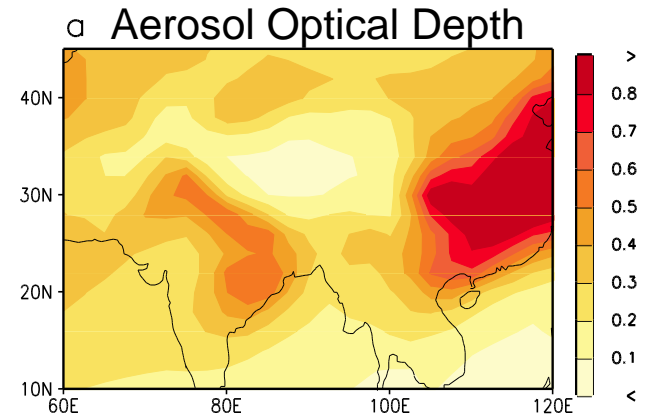


# Effect of absorbing aerosols on snow melt over the Tibetan Plateau (EHP Effect)

(GEOS4-fvGCM with Slab Ocean Model)



1. Slow melting phase (I) in April is initiated by sensible heat transfer from warmer atmosphere to land.
2. Rapid melting phase (IIA) is due to an evaporation-snow-land feedback coupled to an increase in moisture over TP
3. Melting is slowed down by increasing surface evaporation in Phase IIB.
4. Partial recovery (Phase III)

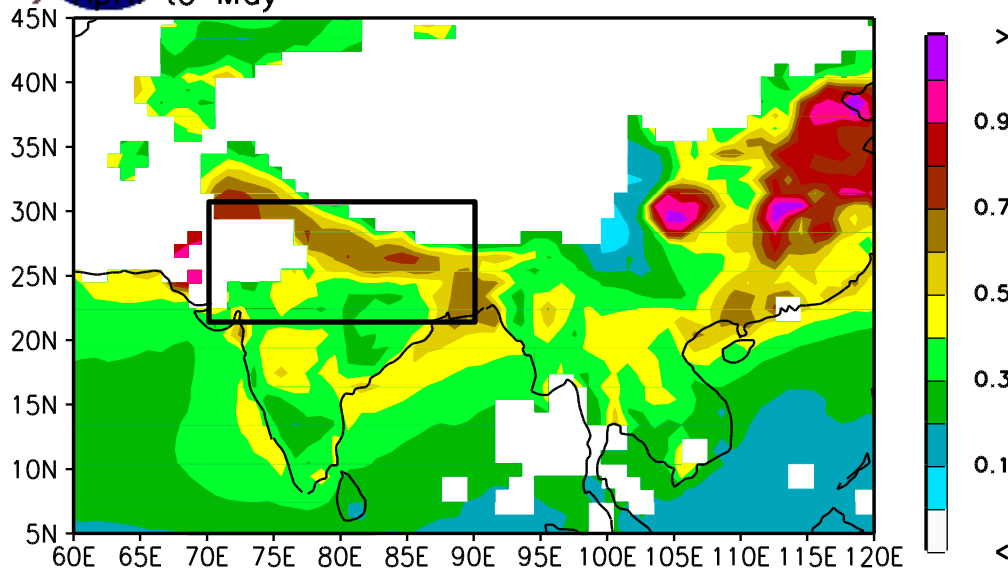


Lau, K. M., M. K. Kim, K. M. Kim, and W. S. Lee, 2010: Enhanced surface warming and accelerated snow melt in the Himalayas and Tibetan Plateau induced by absorbing aerosols. *Environ. Res. Lett.*



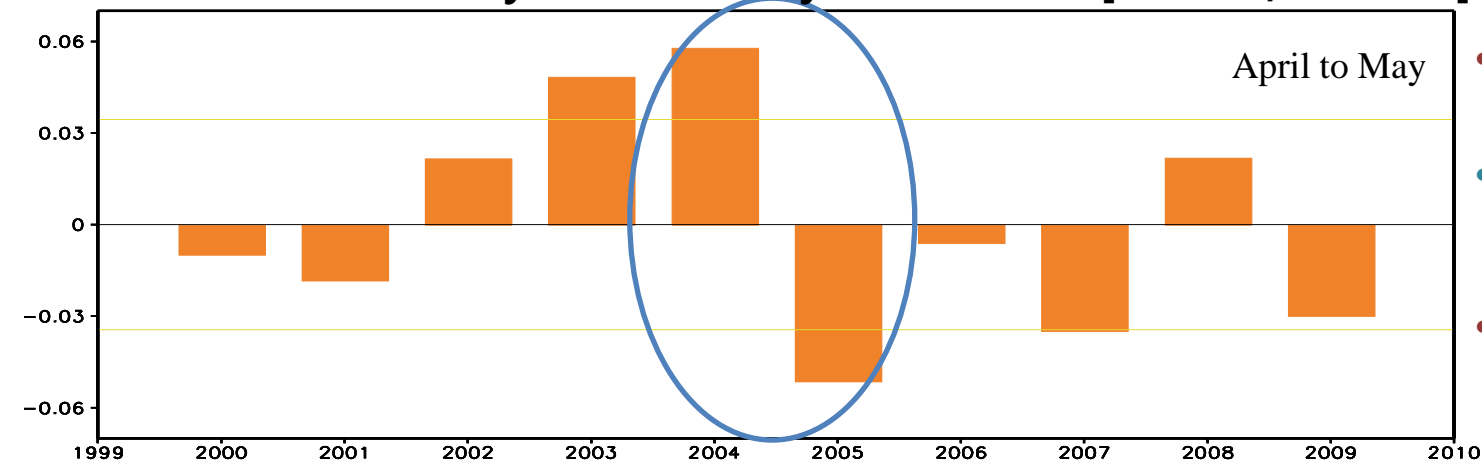
# April-May mean aerosol optical depth(AOD) from 2000 to 2009

April to May



The dusts in IGP are coated with black carbon produced from local emissions and become a strong absorber of solar radiation and an efficient source of atmosphere heating( Lau and Kim, 2006).

Area mean monthly AOD anomaly over IGP [70~90E, 20~30N]



• High AOD : 2004

• Low AOD : 2005

• Nino3 Index

DJF 2004: 0.29

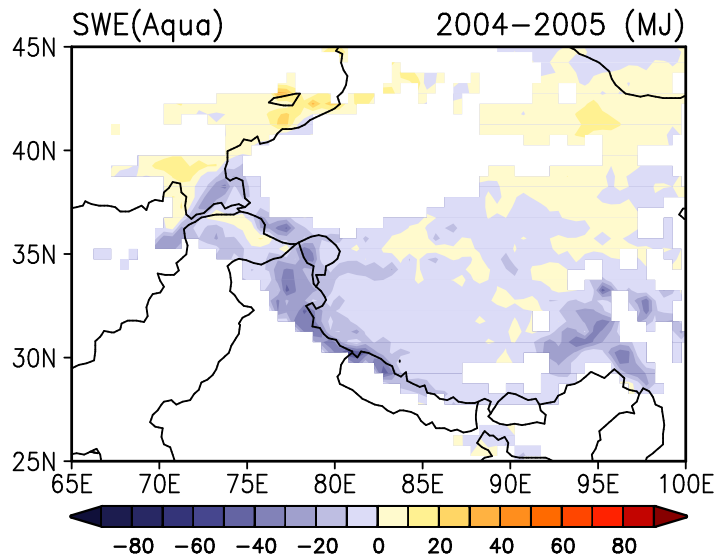
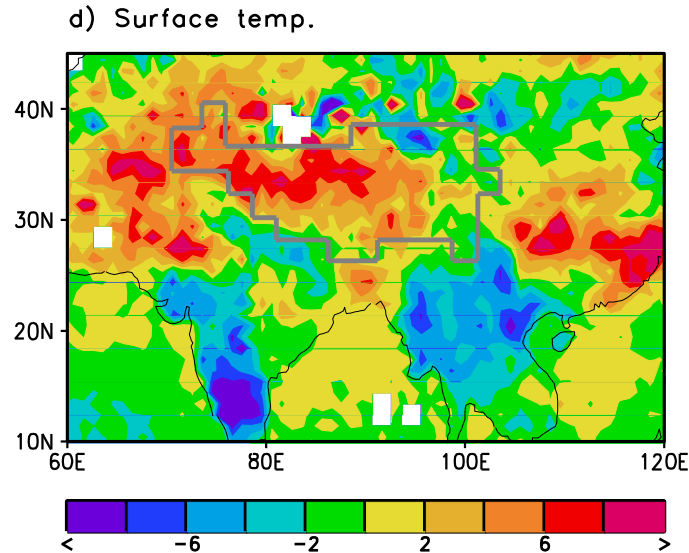
DJF 2005: 0.26

DJF 2003: 0.83



# May/June Difference between 2004 and 2005

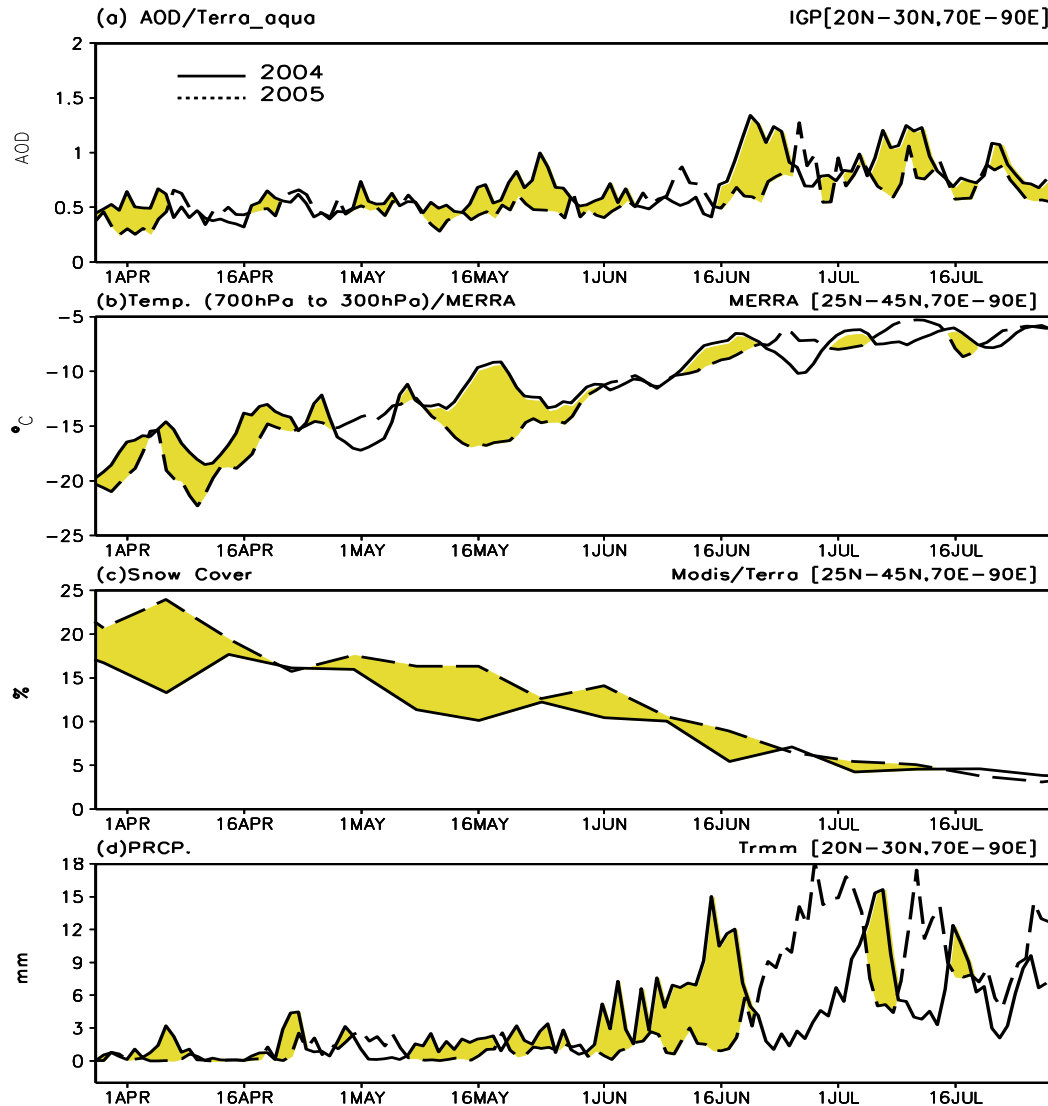
## Dirtier year minus cleaner year





# Intra-seasonal variation

Dirtier year (MJ 2004) minus cleaner year (MJ 2005)



Increased AOD over  
N. India

Increased Upper  
level Temperature  
Tibetan Plateau

Decreased Snow  
Cover over TP

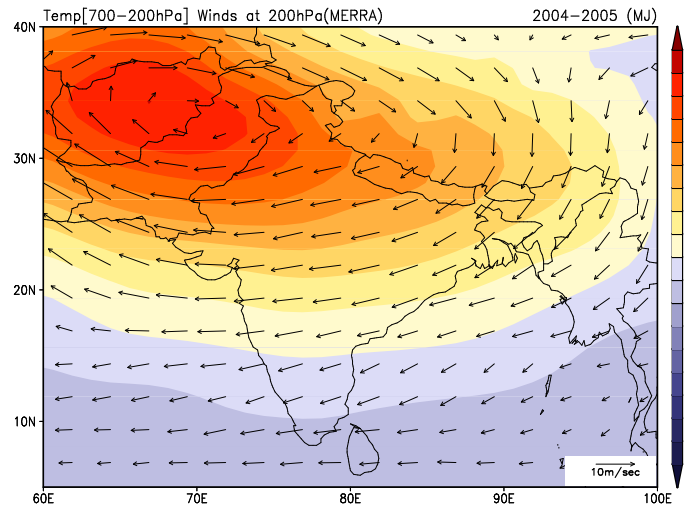
Increased Rainfall  
over N. India



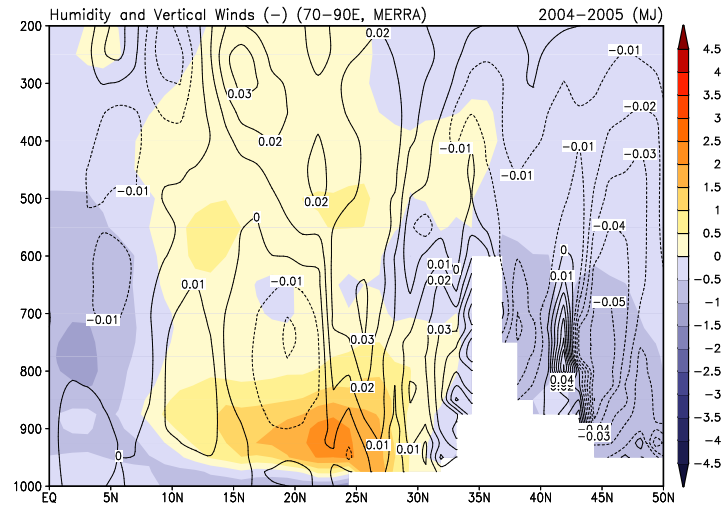
# May/June Difference between 2004 and 2005

## Dirtier year minus cleaner year

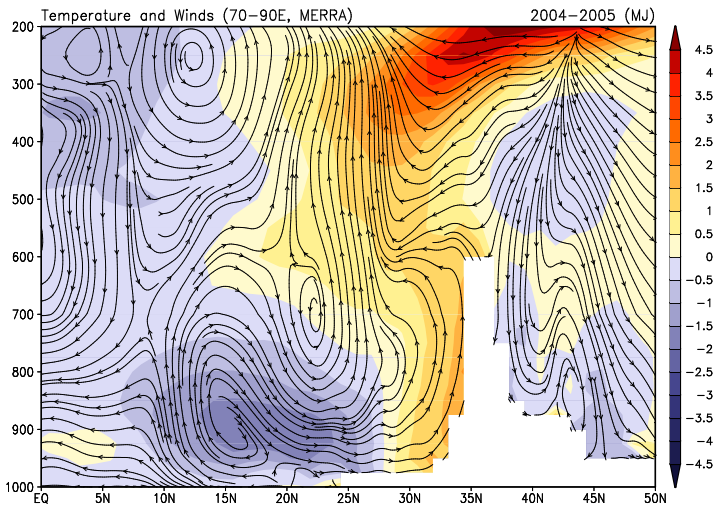
### Upper level Temperature & Wind



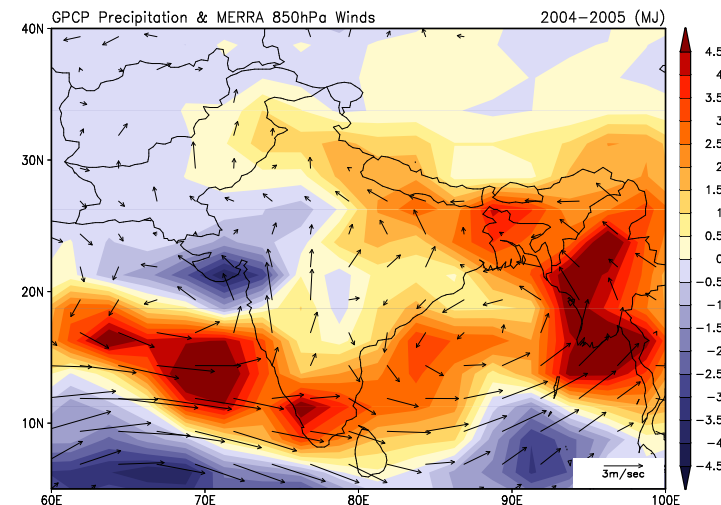
### Moisture & Vertical Winds



### Temperature & winds (v-w)

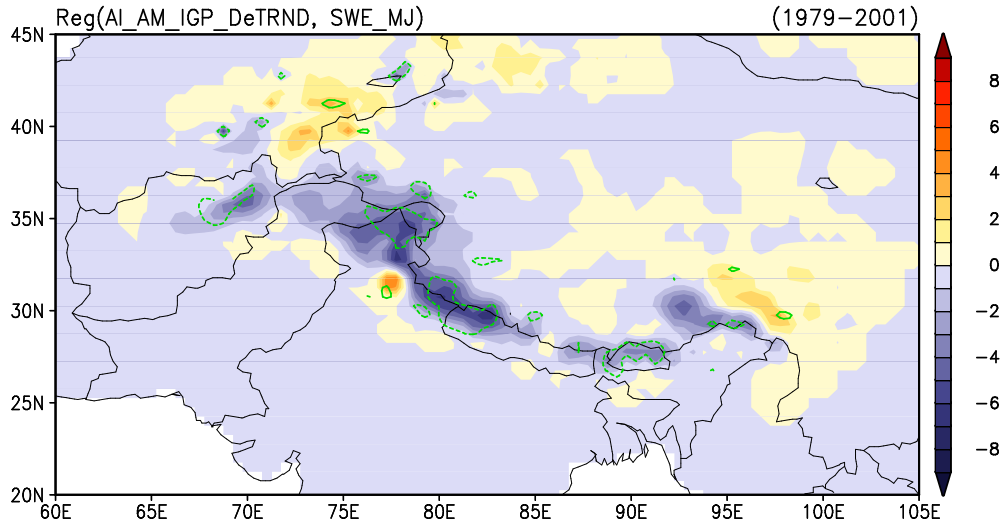


### Lower level Winds & Rainfall





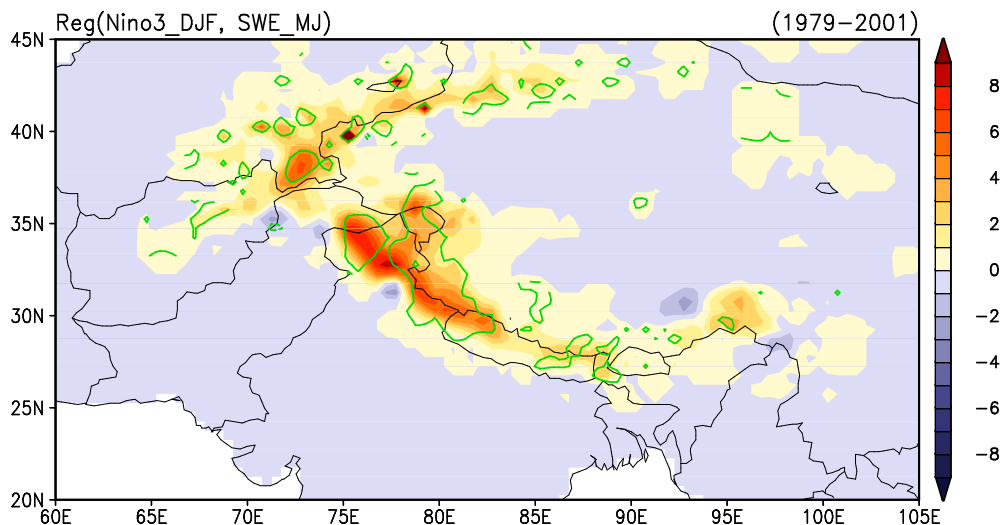
# Interannual Variation of Snow Water Equivalent (Aerosol vs El Nino)



## Aerosol Effect

Heating of the atmosphere by dust and BC can lead to widespread enhanced warming over the TP and accelerated snow melt in the western TP and Himalayas. (Lau et al. 2010)

Decreased snow albedo due to BC/dust deposition can further accelerate snow melt (Yasunari et al. 2010)

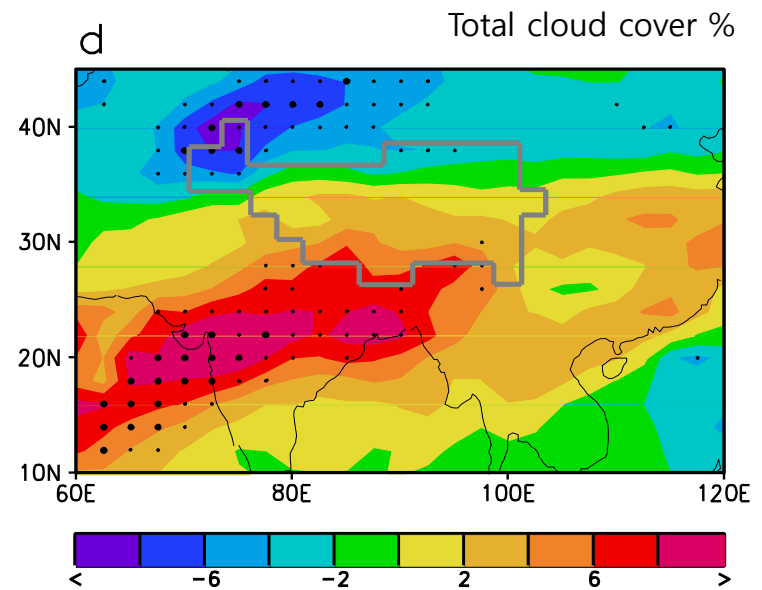
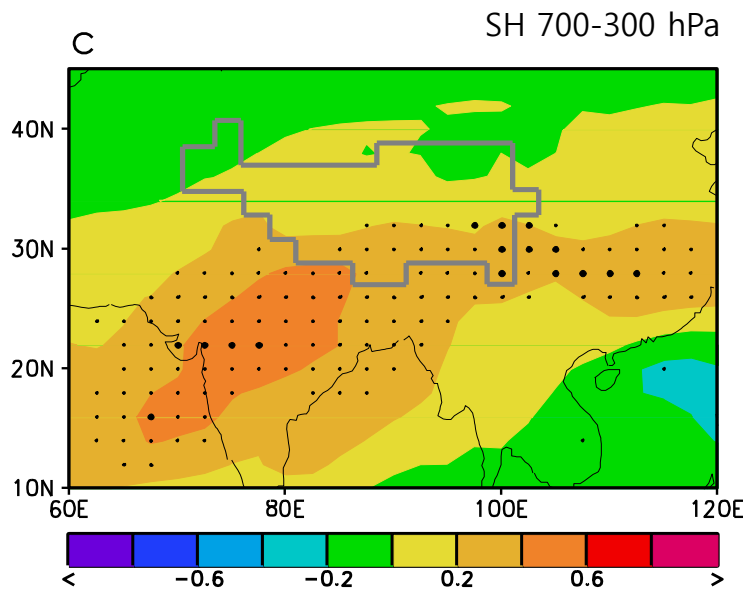
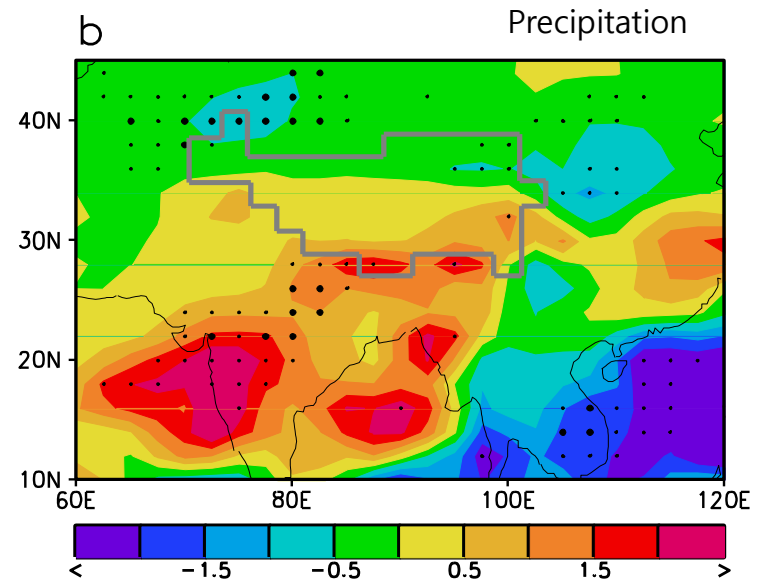
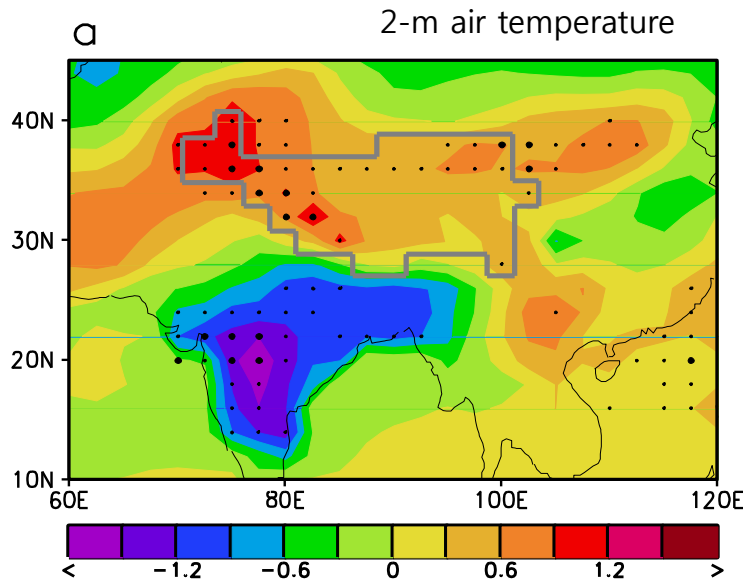


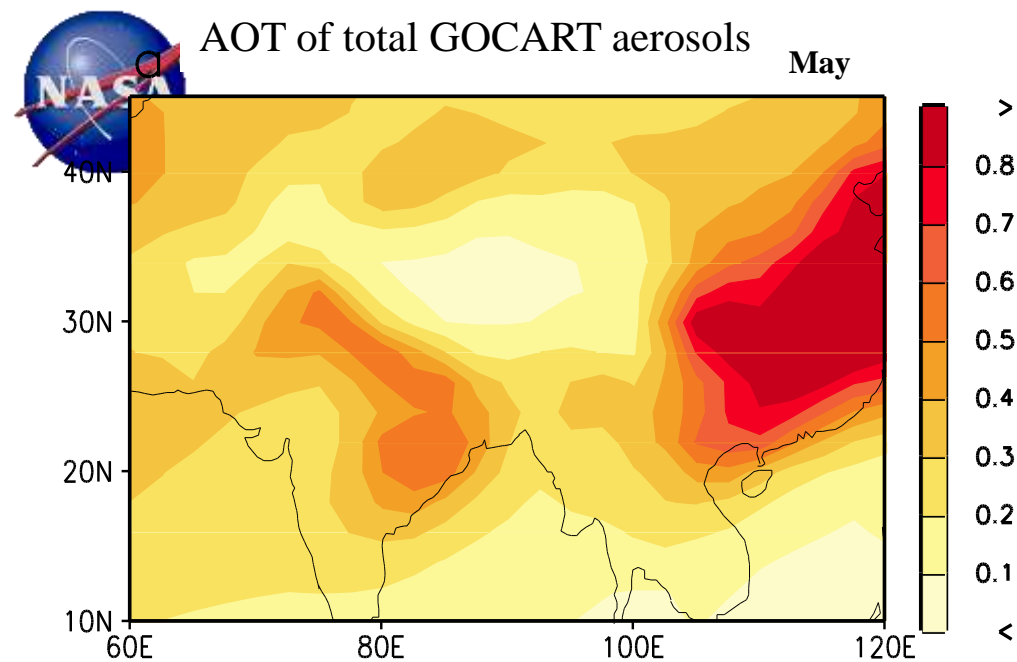
## El Nino Effect

The atmospheric teleconnection pattern initiated by ENSO increase snowfall over TP during winter. The increased snowfall produces a larger snowpack which lasts through the spring and summer, and subsequently weaken the Indian monsoon (Shaman and Tziperman 2005).



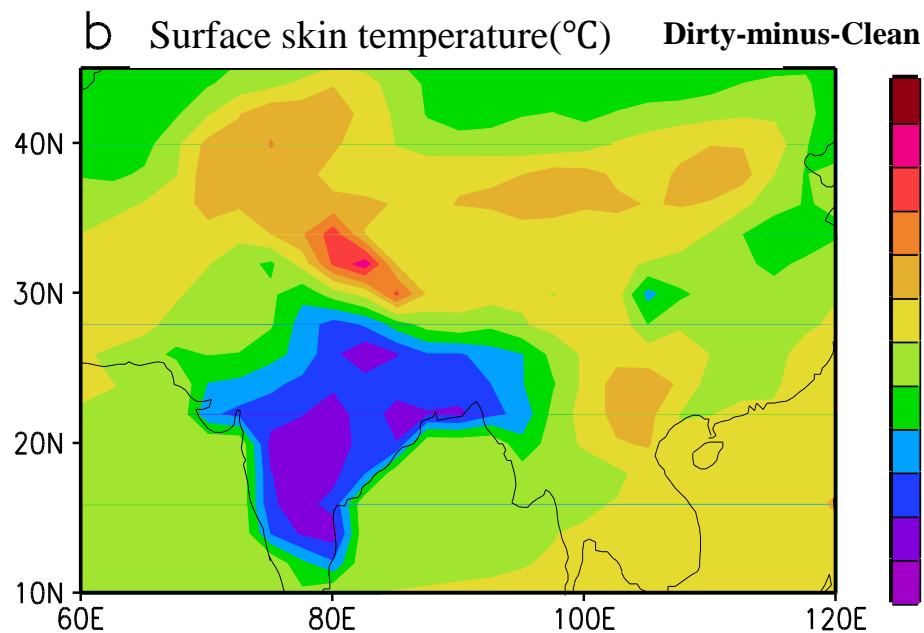
# GEOS GCM, Aerosol-minus-No Aerosol, for month of May





## Area-averaged composition of aerosols in May

Region Aerosol Type	IGP [ 75~90°E, 17.5~30°N]	TP [ 80~95°E, 30~35°N]	Central China [105~120°E, 25~35°N]
All aerosols	0.24	0.15	0.73
Black Carbon	(11.7%)	(10.7%)	( 8.0%)
Dust	(41.9%)	(43.3%)	(15.8%)
Sulfate	(15.8%)	(22.3%)	(62.5%)
Organic Carbon	(26.5%)	(20.4%)	(12.4%)
Sea salt	( 3.8%)	( 3.0%)	( 1.1%)

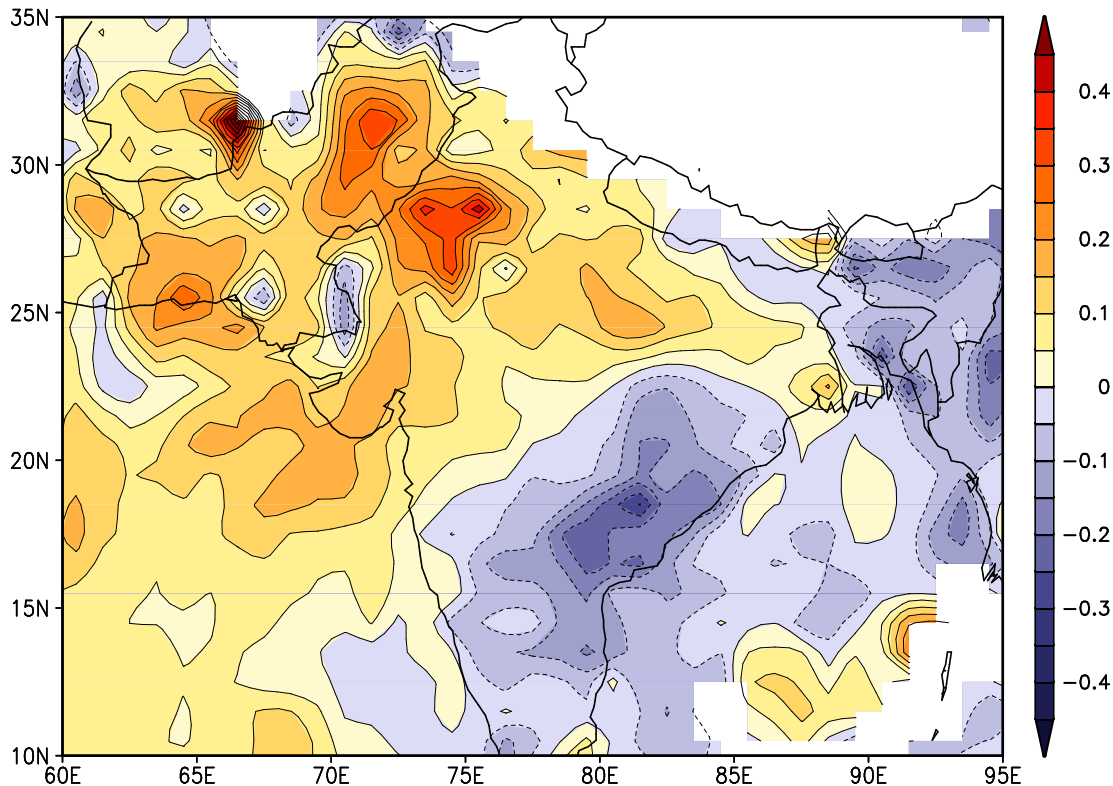




# Observational Case Study

## 2004 vs 2005

Aqua AOD + DeepBlue  
Apr-May (2004 minus 2005)



### Data Used

MODIS AOD

Snow cover (MODIS/Terra)

Snow Water Equivalent  
(SSMR & SSM/I) & (AMSR-E)

MERRA Winds & Temperature

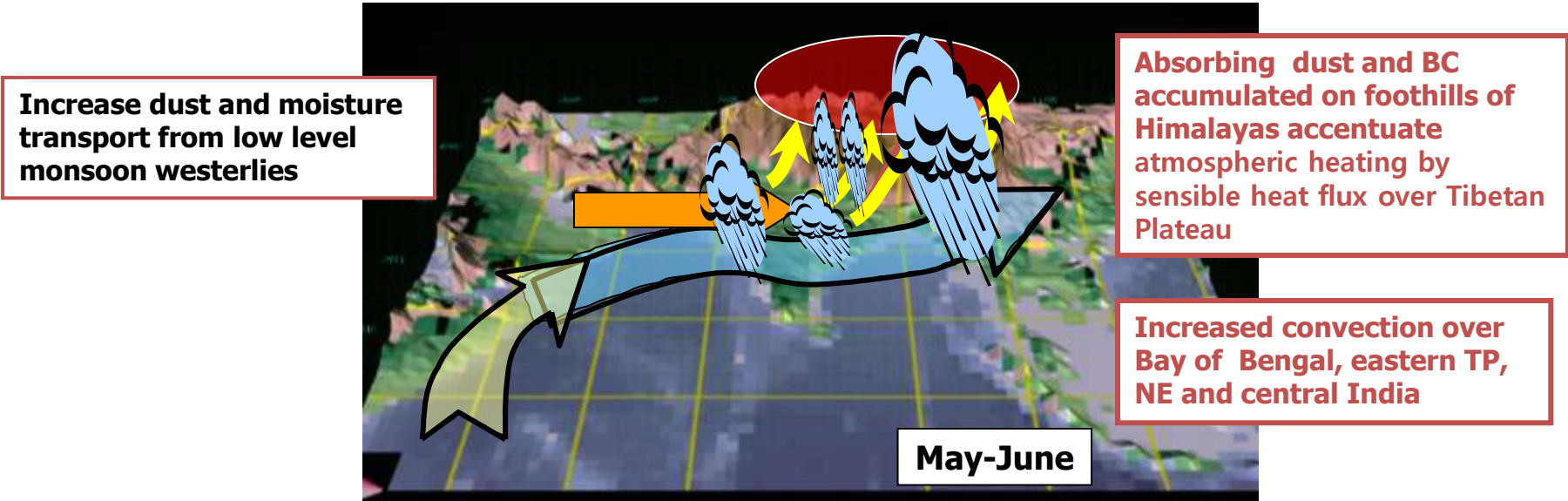
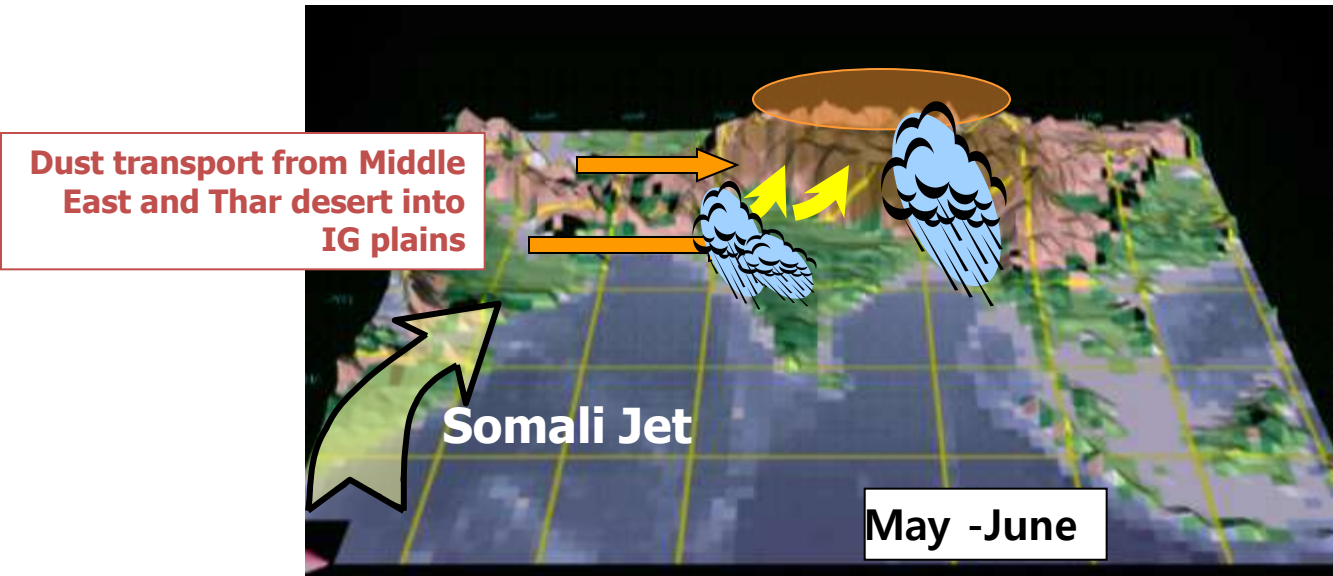
Surface Temperature (AIRS)

Nino3 SST Index



# The Elevated Heat Pump Hypothesis

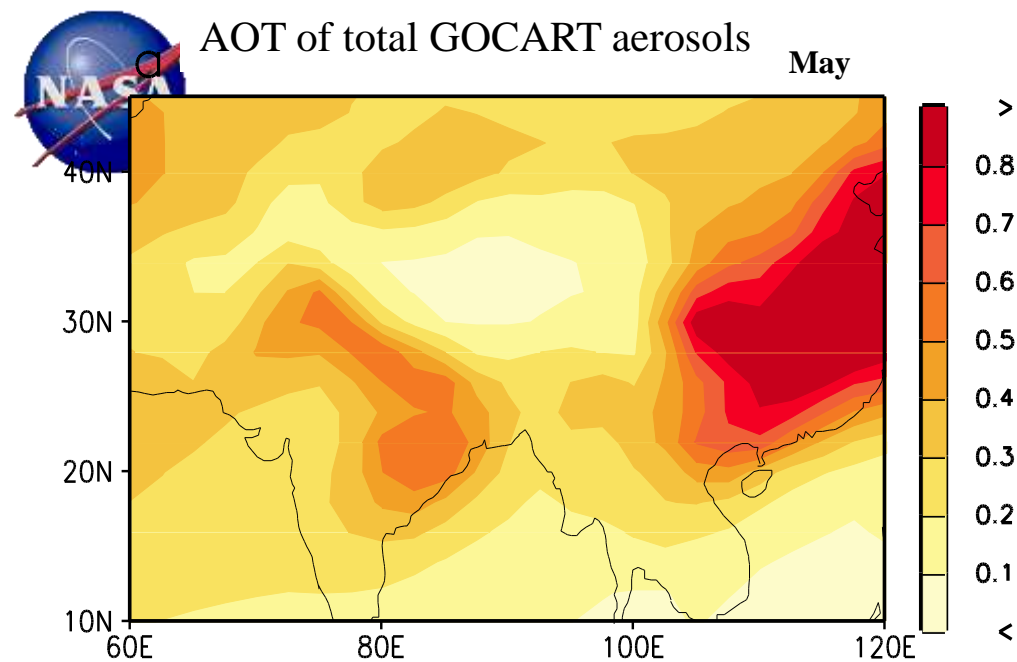
(Lau et al. 2006, Lau and Kim 2006)





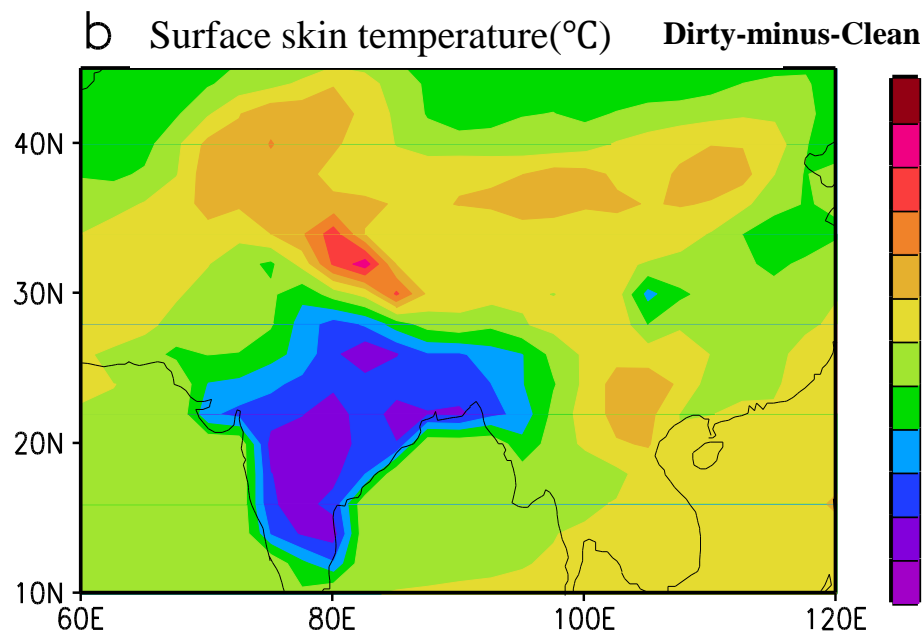
## Possible effects by aerosols on the large-scale monsoon water cycle

- **The Solar Dimming Effect** : More aerosol cooling over land and ocean to the north reduce meridional thermal contrast, causes a spin-down of the Hadley circulation, weakens the monsoon (Ramanathan et al 2005, Chung and Ramanathan, 2005..)
- **Atmospheric heating feedback**: Dust and black carbon piling up against the foothills of the Himalayas, heats up upper troposphere, producing a positive water cycle feedback, enhancing the monsoon (Lau et al. 2006, Lau and Kim, 2006, 2007, Meehl et al. 2007, Collier and Zhang 2009, Wang 2009...)
- **Microphysics effects of aerosol** – a wild card
- **Coupled Atmosphere-land-ocean dynamics**: All aerosols effects are likely to be confounded /modulated by large-scale remote forcing, e.g., El Nino, local SST, global warming...



## Area-averaged composition of aerosols in May

Region Aerosol Type	IGP [ 75~90°E, 17.5~30°N]	TP [ 80~95°E, 30~35°N]	Central China [105~120°E, 25~35°N]
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Sulfate	(15.8%)	(22.3%)	(62.5%)
Organic Carbon	(26.5%)	(20.4%)	(12.4%)
Sea salt	( 3.8%)	( 3.0%)	( 1.1%)





# OUTLINE

- Asian Climate Change primer
- Possible impacts of aerosols on
  - Asian monsoon water cycle
  - Accelerated melting of snowpack in the Himalayas and Tibetan Plateau
- Some thoughts on adaptation and mitigation

# The Greenhouse Effect

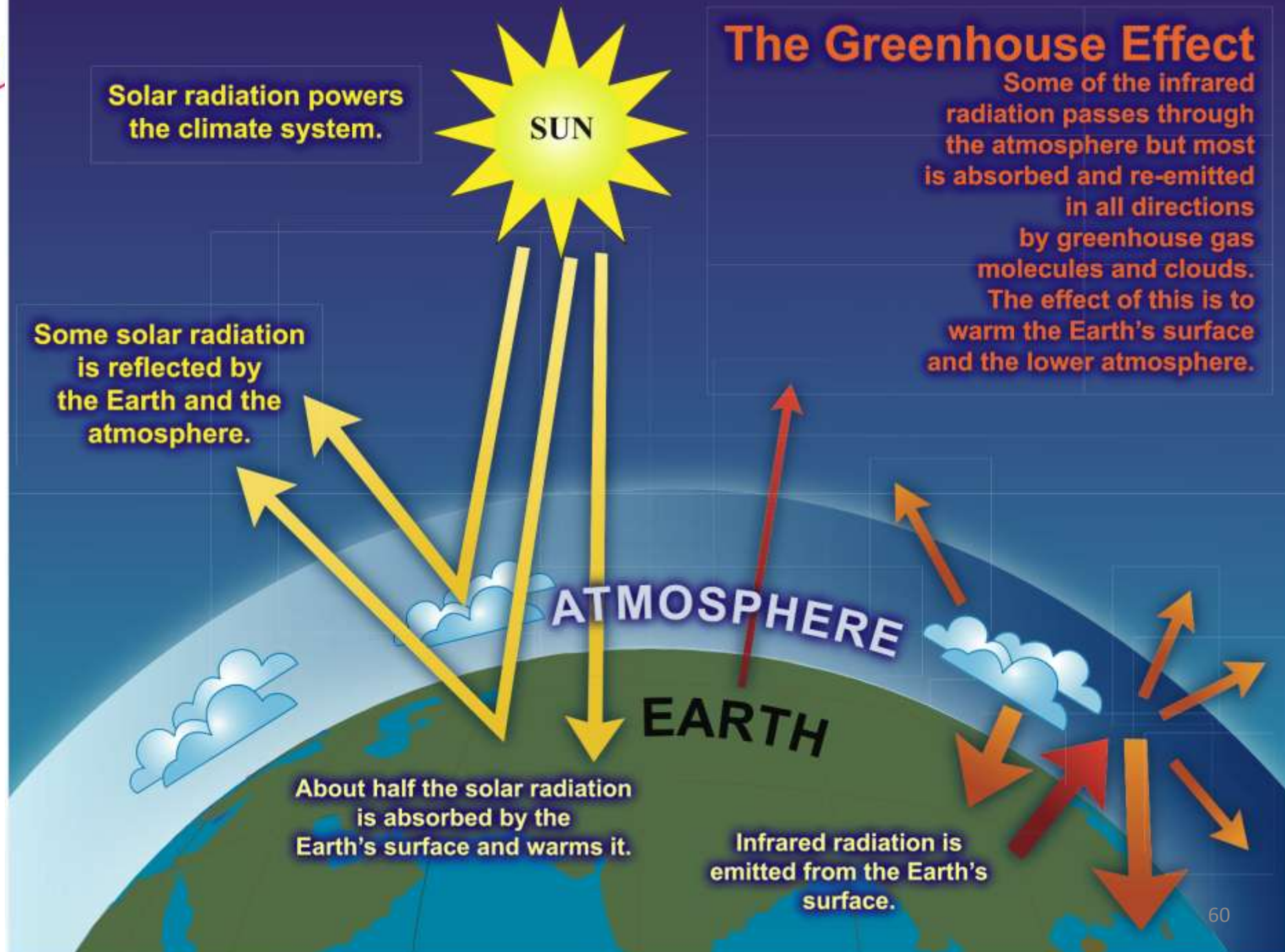
Some of the infrared radiation passes through the atmosphere but most is absorbed and re-emitted in all directions by greenhouse gas molecules and clouds. The effect of this is to warm the Earth's surface and the lower atmosphere.

Solar radiation powers the climate system.

Some solar radiation is reflected by the Earth and the atmosphere.

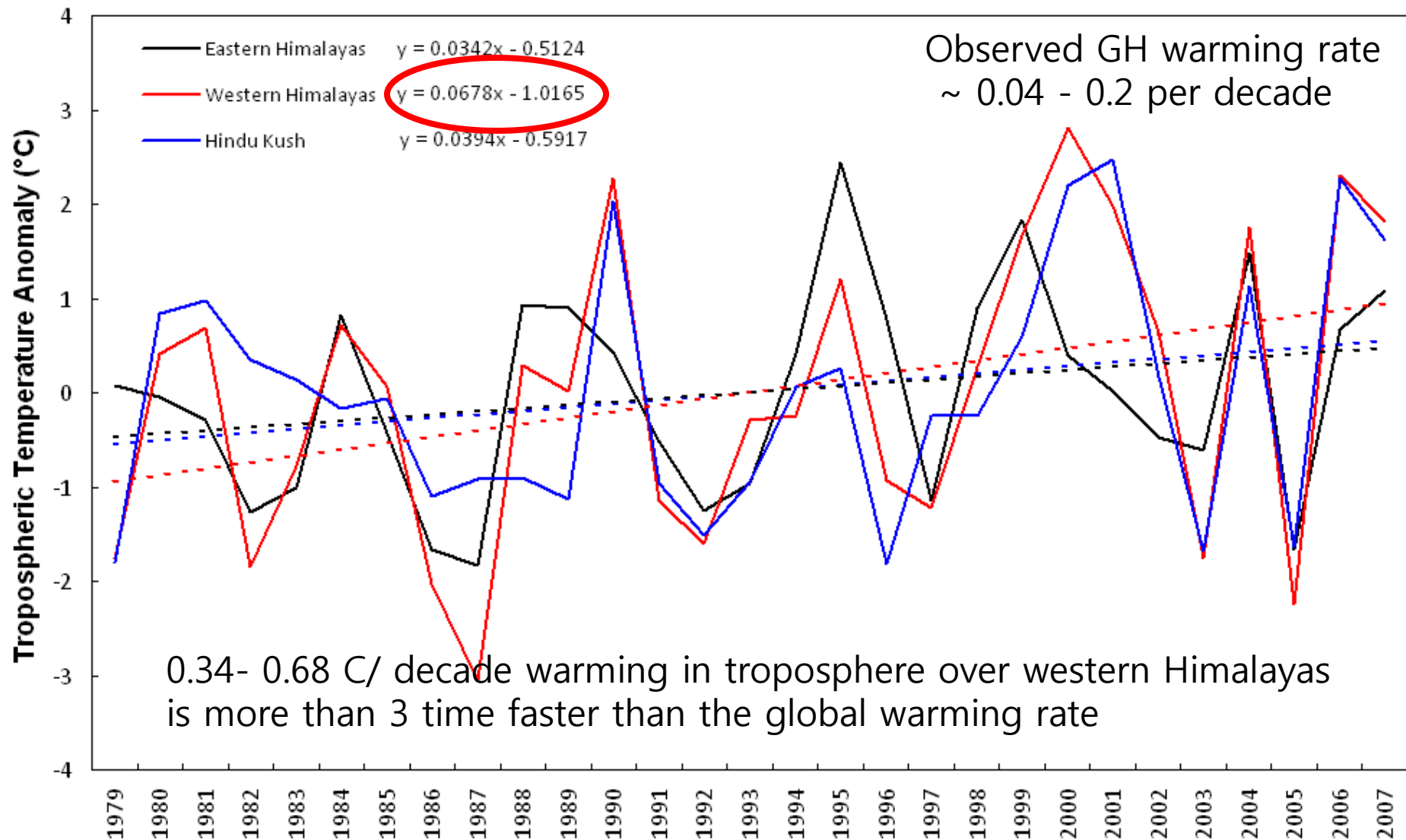
About half the solar radiation is absorbed by the Earth's surface and warms it.

Infrared radiation is emitted from the Earth's surface.



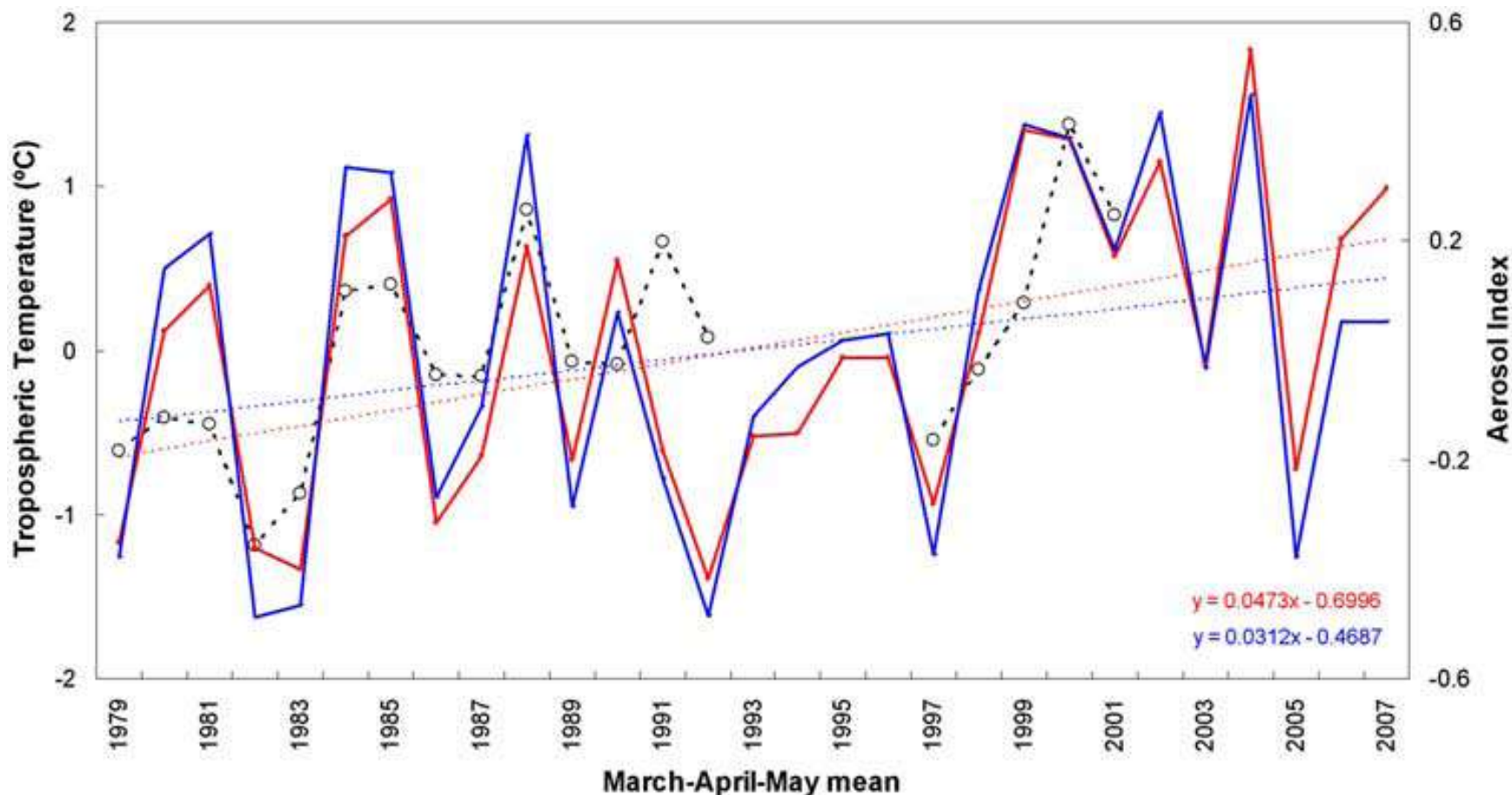


## May (1979-2007)





# Inter-annual variability of pre-monsoon Tropospheric Temperatures and Aerosol loading



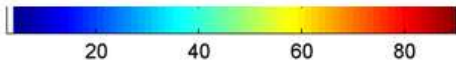
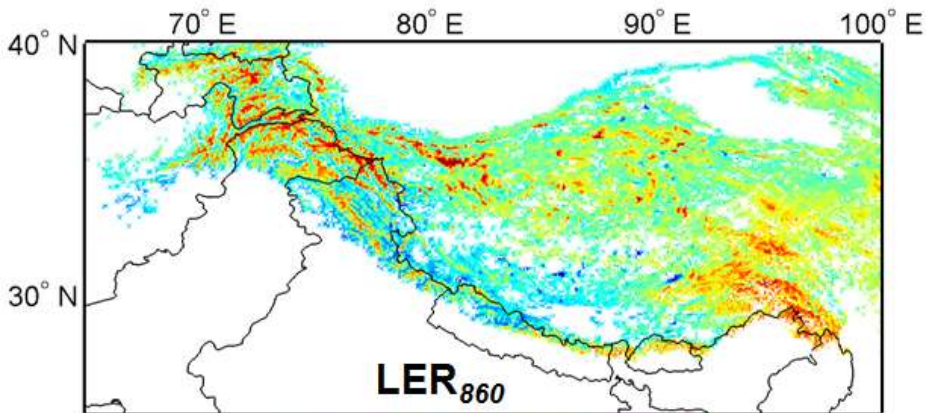
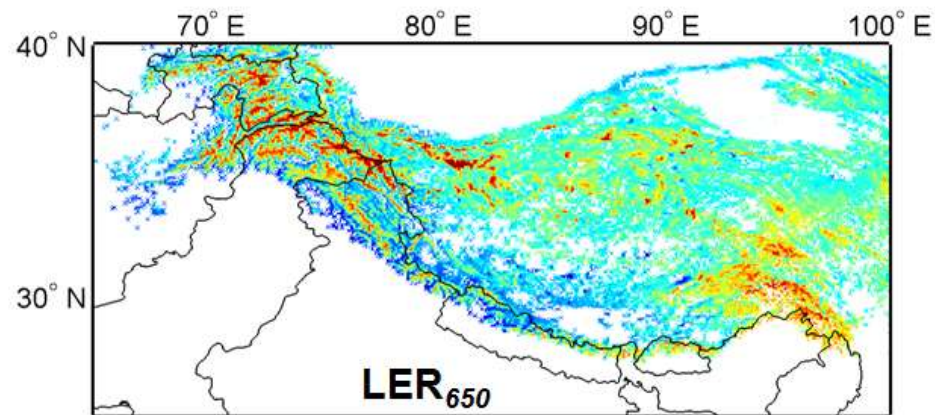
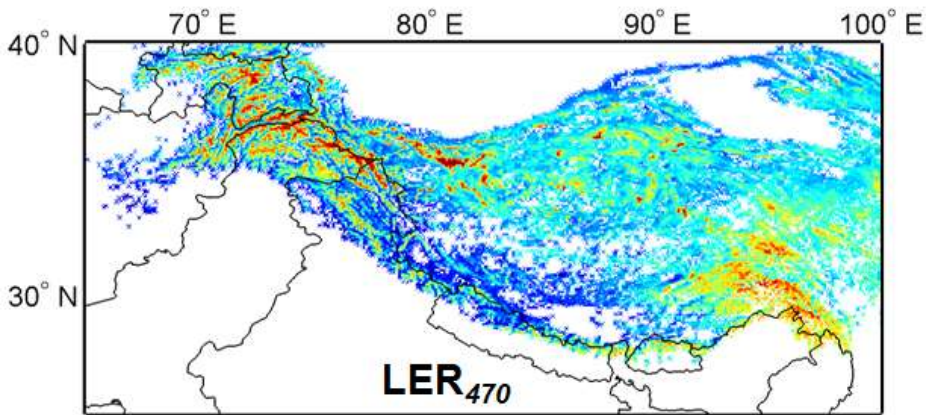
- Mid-Tropospheric Temp
- Lower Tropospheric Temp
- - - TOMS Aerosol Index

Correlation coef. Of 0.67 and 0.65  
between  $T_{mid}$  and  $T_{low}$  with AI (exceeding 95% significance)



# MODIS-derived Lambertian Equivalent Reflectance (LER) over the Himalayan-Tibetan Plateau Snow Surface at TOA

May 15 – June 15 LER





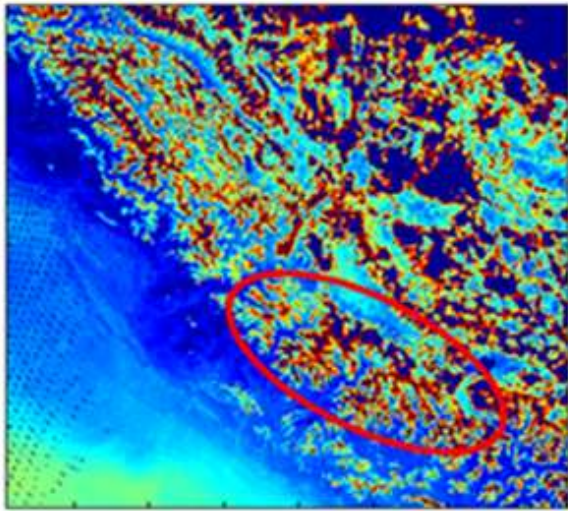
The spectral reflectance from MODIS suggests perturbation caused due to the dust layer over the snow surface as well as the reduced reflectance associated with dust deposition over snow covered pixels in the satellite scene

6 June 2003

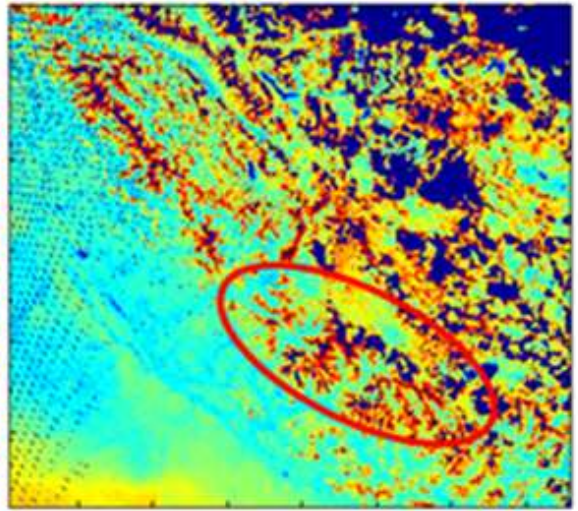
**Terra MODIS RGB**



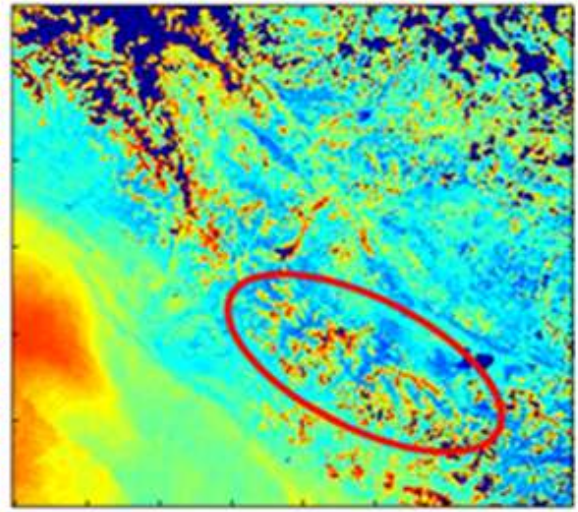
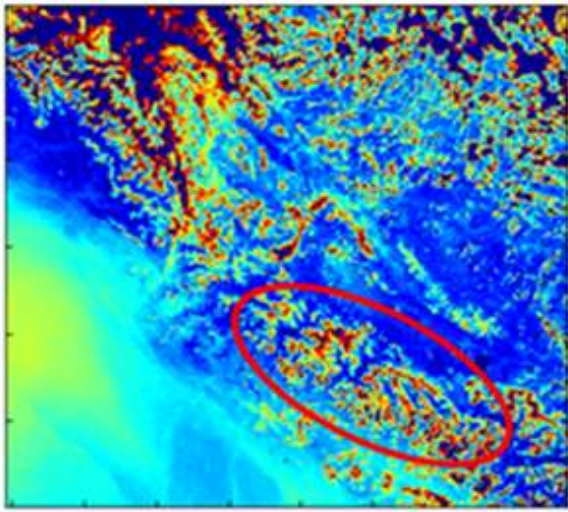
**LER<sub>470</sub>**



**LER<sub>860</sub>**

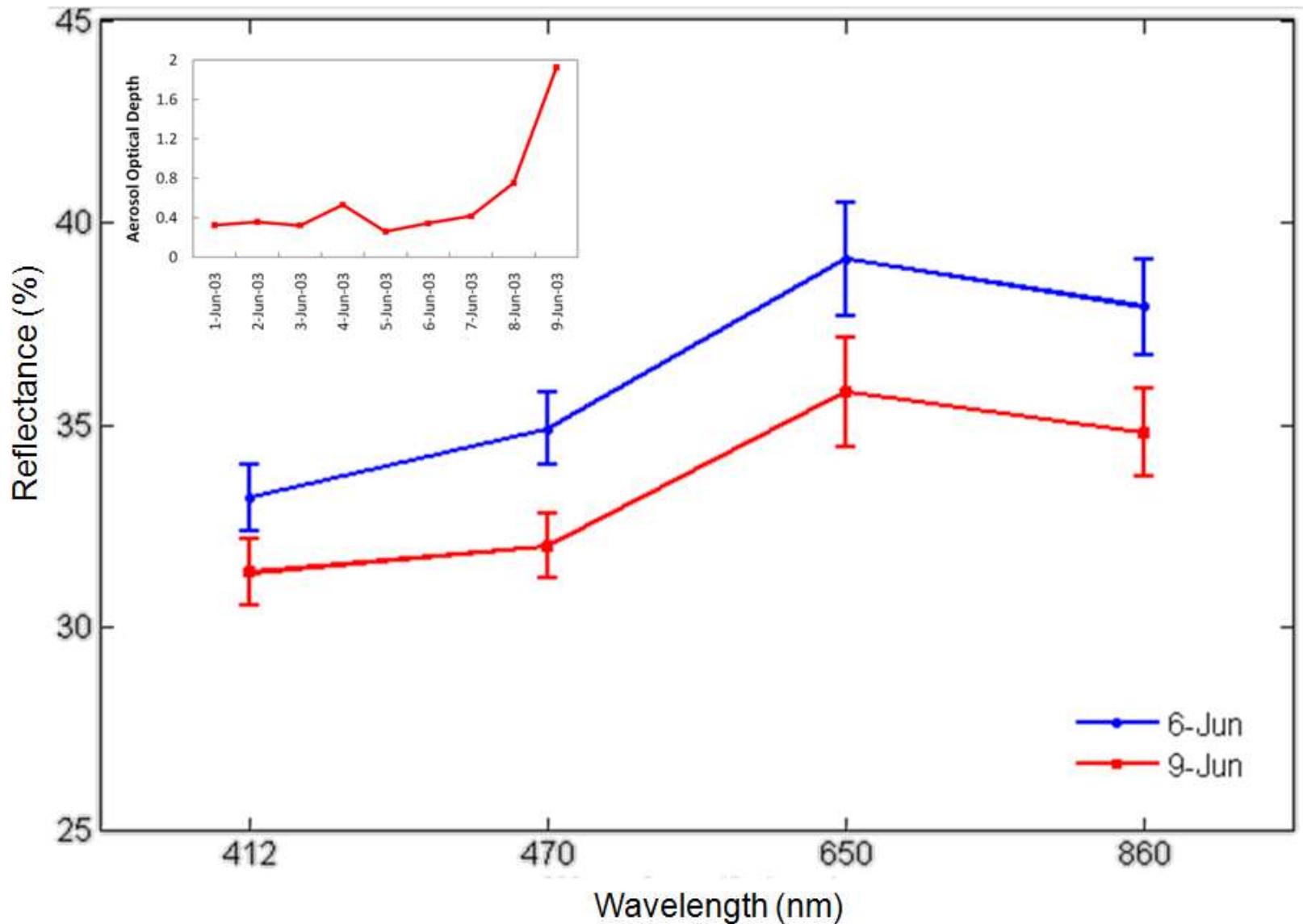


9 June 2003





Reduction of 5-8% in LER during dust transport. Inset plot shows anomalous high AOD on 9 June 2003 over foothills associated with the extensive dust plume.





Preliminary estimates, based on 14 days of Terra/Aqua MODIS observations during pre-monsoon season, of dust overlaying western Himalayan snow surface, indicate characteristic changes in spectral reflectance with 3 - 6% reduction at shorter wavelengths (412 and 470nm), while 5 - 7% decrease at 650 and 860nm channels.

